Introduction to 2D Hydraulics Equations

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Hydraulic Modeling

Formal Framework:
- Mass Conservation
- Momentum Conservation
- Empirical Equations

Computational Concepts:
- Computational Mesh
- Numerical Methods
- Subgrid Bathymetry

HEC-RAS

Data:
- Topography, Maps, LIDAR
- Historical Information
- Modeler Experience

Engineering Problems:
- Project Planning & Design
- Dam and Levee Breaks
- Overland Flows, Floods

Outline

- Mass Conservation (Continuity)
- Momentum Conservation (Depth-Averaged)
  - Acceleration
  - Coriolis term
  - Hydrostatic pressure
  - Turbulent mixing
  - Friction
- Diffusion Wave Equation
- Numerical Methods

Review of Basic Concepts

- Continuity: Fluid mass cannot be created nor destroyed
- Uniform Flow: Flow is uniform in space
- Steady Flow: Flow is uniform in time
- Laminar Flow: Translational flow (high viscous force relative to inertial force)
- Turbulent Flow: Chaotic flow (small viscous force relative to inertial force)
- Shear stress: Tangential force per unit area
- Pressure: Normal force per unit area
- Streamline: Line drawn through flow where every point is tangential to velocity vector

Introduction

- Shallow Water Equations
  - **System** of partial differential equations with **many forms** which arise the simulation of fluid flow in **rivers**, oceans, coastal regions, atmospheric flows, and **debris** flows
  - **Main assumption:** vertical accelerations much smaller than horizontal accelerations
  - **Derived from:** Navier-Stokes equations which describe the conservation of mass and linear momentum in fluids

Mass Conservation

- Assuming a constant water density

$$\frac{\partial h}{\partial t} + \nabla \cdot (h\boldsymbol{V}) = q$$

- Integrating over a computational cell

$$\frac{\partial}{\partial t}\iiint_{\Omega} d\Omega + \iint_{S} (\boldsymbol{V} \cdot \boldsymbol{n}) dS = Q$$

- Finite-Volume Discretization

$$\frac{\Omega_i^{n+1} - \Omega_i^n}{\Delta t} + \sum_{k \in \mathrm{K}(i)} (\boldsymbol{V}_k \cdot \boldsymbol{n}_{ik}) A_k = Q_i$$

h : Water depth

q : Water souce/sink

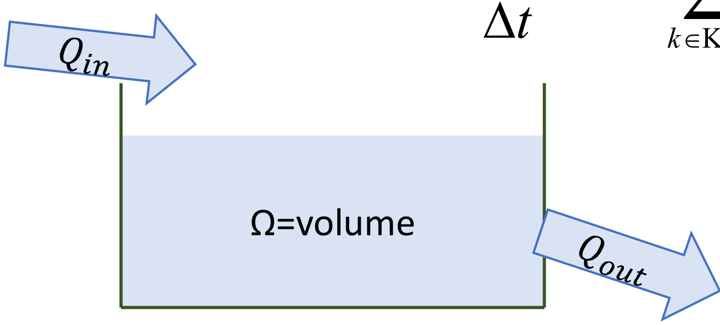
Ω_i : Cell water volume

A_k : Face area

$\boldsymbol{V}_k$: Face velocity

$\boldsymbol{n}_{ik}$: Outward face-normal unit vector

Δt : Time step



Change in volume in a system balances with flow through boundaries

Momentum Conservation

- Momentum Equation (non-conservative form)

$$\underbrace{\frac{\partial \boldsymbol{V}}{\partial t}}_{\text{Temporal}} + \underbrace{(\boldsymbol{V}\cdot\nabla)\boldsymbol{V}}_{\text{Advection}} + \underbrace{f_c \boldsymbol{k}\times\boldsymbol{V}}_{\text{Coriolis}} = -\underbrace{g\nabla z_s}_{\text{Pressure gradient}} + \underbrace{\frac{1}{h}\nabla\cdot\left(\boldsymbol{v}_t h\nabla \boldsymbol{V}\right)}_{\text{Diffusion}} - \underbrace{\frac{\boldsymbol{\tau}_b}{\rho R}}_{\text{Bottom Friction}} + \underbrace{\frac{\boldsymbol{\tau}_s}{\rho h}}_{\text{Wind Stress}}$$

- From Newton's 2nd Law of motion (i.e. F=ma)
- Momentum: M = mV
- Assumes constant water density, small vertical velocities, hydrostatic pressure, etc.
- Non-linear and a function of both velocity and water levels
- Continuity and Momentum Equations are the Shallow Water Equations or sometimes referred to as the "Full Momentum" equations in HEC-RAS

$\boldsymbol{V}$: Velocity

z_s : Water level

g : Gravity

$\boldsymbol{v}_t$: Turbulent eddy viscosity

h : Water depth

R : Hydraulic Radius

f_c : Coriolis Parameter

τ_b : Bed shear stress

τ_s : Surface stress

Accelerations

- **Eulerian:** Frame of reference fixed in space and time

$$\frac{\partial \boldsymbol{V}}{\partial t} + (\boldsymbol{V} \cdot \nabla)\boldsymbol{V}$$

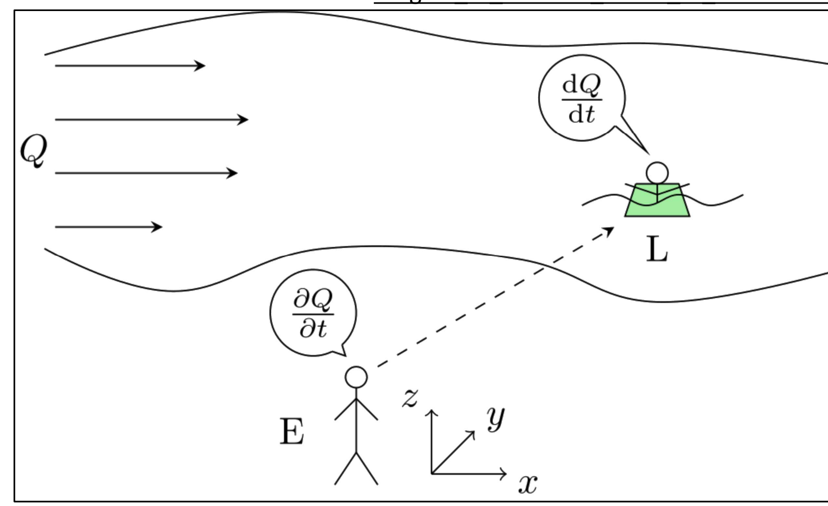
  - Easier to compute
  - Time-step restricted by Courant condition

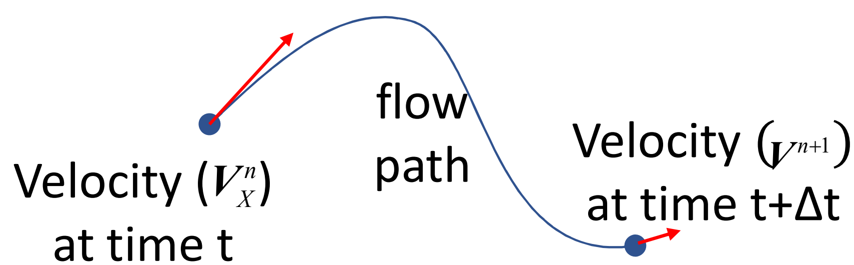
- **Lagrangian:** Frame of reference moves with total derivative along flow path

$$\frac{\partial \boldsymbol{V}}{\partial t} + (\boldsymbol{V} \cdot \nabla)\boldsymbol{V} = \frac{D\boldsymbol{V}}{Dt} = \frac{\boldsymbol{V}^{n+1} - \boldsymbol{V}_X^{n}}{\Delta t}$$

  - More expensive to compute
  - Allows larger time-steps

https://commons.wikimedia.org/wiki/File:Lagrangian_vs_Eulerian_frame_of_reference.svg



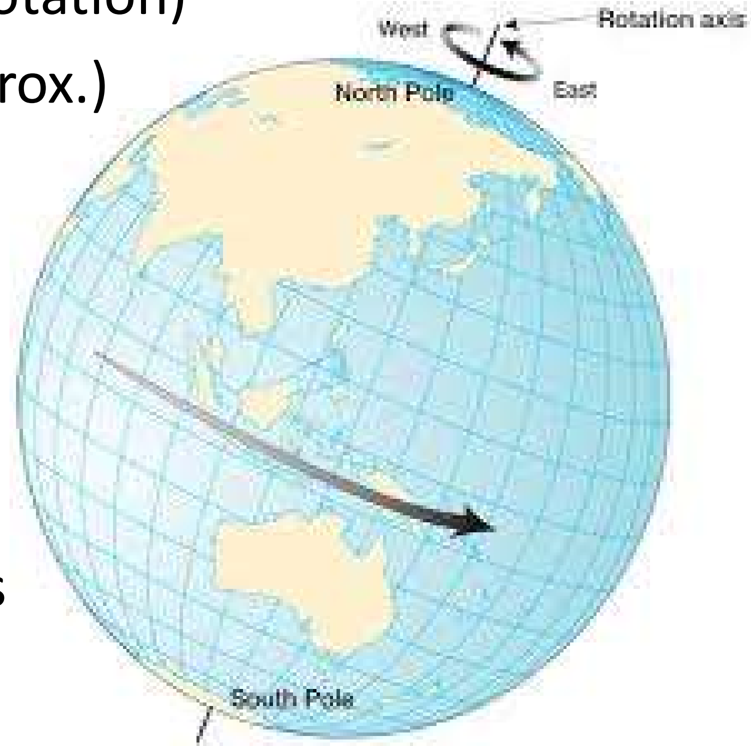


Coriolis Acceleration

- Effect of rotating frame of reference (earth's rotation)
- Constant for the each 2D domain (f-plane approx.)

$$f_c = 2\,\omega \sin\varphi$$

  - ω: sidereal angular velocity of the Earth
  - φ: latitude. Positive for northern hemisphere. Negative for southern hemisphere
- Coriolis acceleration disabled by default to save computational time
- Negligible for most river and flood simulations
- When to enable Coriolis term?
  - Large domains
  - Higher latitudes



Pressure Gradient

- Assumes vertical water accelerations are small compared to gravity
- Total pressure is

$$P = P_{atm} + \rho g(z_s - z)$$

  - P_{atm}: atmospheric pressure (assumed to be constant)
  - ρ : constant water density
  - g: gravity acceleration constant
  - z_s: water surface elevation
  - z: vertical coordinate
- Pressure gradient

$$\nabla P = \nabla P_{atm} + \rho g \nabla z_s$$

Bottom Friction

- Resisting force due to relative motion of fluid against the bed
- Bed Shear Stress

$$\boldsymbol{\tau}_b = \rho C_b \left|\boldsymbol{V}\right| \boldsymbol{V}$$

n :Manning coefficient

ρ: water density

g: gravity acceleration constant

$|\boldsymbol{V}|$: velocity magnitude

R: hydraulic radius

- Drag Coefficient

$$C_b = \frac{gn^2}{R^{1/3}}$$

- Friction coefficient

$$c_f = \frac{C_b}{R}\left|\boldsymbol{V}\right| = \frac{gn^2}{R^{4/3}}\left|\boldsymbol{V}\right|$$

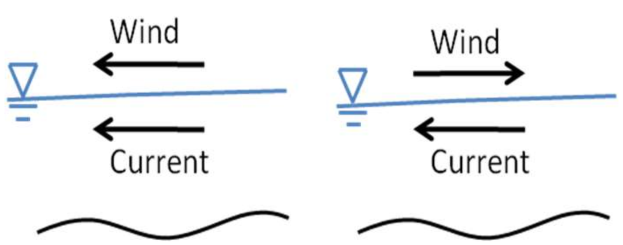
Wind Stress

- Surface Stress is given by

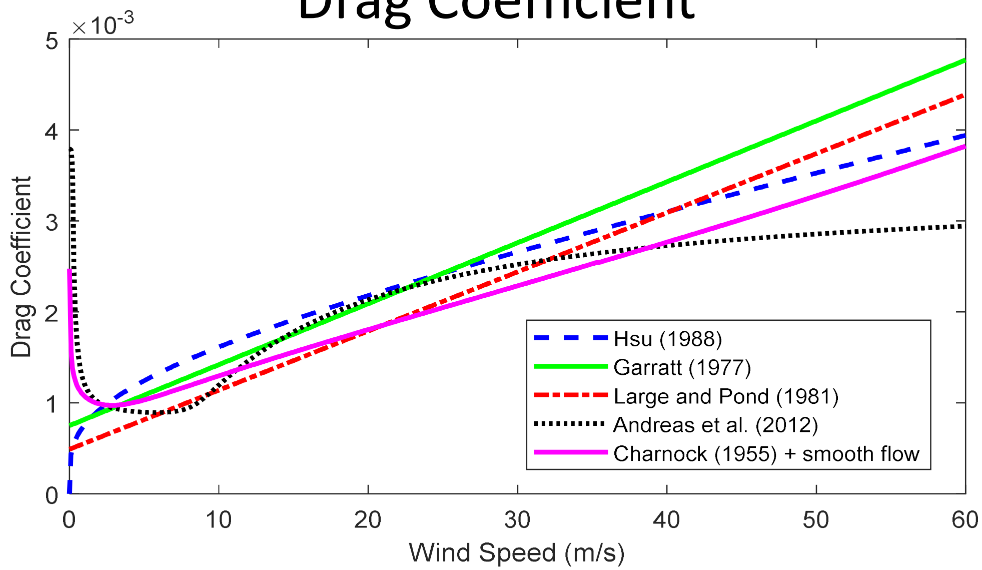
$$\boldsymbol{\tau}_s = \rho_a C_D \left|\boldsymbol{W}_{10}\right| \boldsymbol{W}_{10}$$

- Wind Reference Frame

$$\boldsymbol{W}_{10} = \begin{cases} \boldsymbol{W}_{10}^E - \boldsymbol{V} & \text{for Lagrangian} \\ \boldsymbol{W}_{10}^E & \text{for Eulerian} \end{cases}$$



Drag Coefficient



Diffusive-Wave Approximation

- Ignoring the following terms

$$\underbrace{\cancel{\frac{\partial \boldsymbol{V}}{\partial t}}}_{\text{Temporal}} + \underbrace{\cancel{(\boldsymbol{V}\cdot\nabla)\boldsymbol{V}}}_{\text{Advection}} + \underbrace{\cancel{f_c \boldsymbol{k}\times\boldsymbol{V}}}_{\text{Coriolis}} = \underbrace{-g\nabla z_s}_{\text{Water surface gradient}} + \underbrace{\cancel{\frac{1}{h}\nabla\cdot(\boldsymbol{v}_t h\nabla \boldsymbol{V})}}_{\text{Diffusion}} - \underbrace{\frac{\boldsymbol{\tau}_b}{\rho R}}_{\text{Bottom Friction}} + \underbrace{\frac{\boldsymbol{\tau}_s}{\rho h}}_{\text{Wind Stress}} + \underbrace{\frac{1}{\rho}\nabla P_{atm}}_{\text{Air Pressure}}$$

- Expanding and dividing both sides by the square of its norm leads to

$$\boldsymbol{V} = -\frac{\beta}{h}\nabla z_s \qquad \beta = \frac{R^{2/3}h}{n}\left|\nabla z_s + \frac{\nabla P_{atm}}{\rho g} - \frac{\boldsymbol{\tau}_s}{\rho g h}\right|^{-1/2}$$

- Inserting the above equation into the Continuity Equation leads to the Diffusion-Wave Equation (DWE)

$$\frac{\partial h}{\partial t} = \nabla\cdot(\beta\nabla z_s) + S + q \qquad S = \nabla\cdot\left[\beta\left(\frac{\nabla P_{atm}}{\rho g} - \frac{\boldsymbol{\tau}_s}{\rho g h}\right)\right]$$

SWE vs. DWE

- Use SWE for:
  - Flows with dynamic changes in acceleration
  - Studies with important wave effects, tidal flows
  - Detail solution of flows around obstacles, bridges or bends
  - Simulations influenced by Coriolis, mixing, or wind
  - To obtain high-resolution and detailed flows
- Use DWE for:
  - Flow is mainly driven by gravity and friction
  - Fluid acceleration is monotonic and smooth, no waves
  - To compute approximate global estimates such as flood extent
  - To assess approximate effects of dam breaks
  - To assess interior areas due to levee breeches
  - For quick estimations or preliminary runs

Local Inertial Approximation

- Ignoring the following terms

$$\underbrace{\frac{\partial \boldsymbol{V}}{\partial t}}_{\text{Temporal}} + \underbrace{\cancel{(\boldsymbol{V}\cdot\nabla)\boldsymbol{V}}}_{\text{Advection}} + \underbrace{\cancel{f_c \boldsymbol{k}\times\boldsymbol{V}}}_{\text{Coriolis}} = \underbrace{-g\nabla z_s}_{\text{Water surface gradient}} + \underbrace{\cancel{\frac{1}{h}\nabla\cdot(\boldsymbol{v}_t h \nabla \boldsymbol{V})}}_{\text{Diffusion}} \underbrace{-\frac{\boldsymbol{\tau}_b}{\rho R}}_{\text{Bottom Friction}} + \underbrace{\frac{\boldsymbol{\tau}_s}{\rho h}}_{\text{Wind Stress}} + \underbrace{\frac{1}{\rho}\nabla P_{atm}}_{\text{Air Pressure}}$$

- Also known as the Gravity-Wave Equations

- Compared to DWE
  - Includes temporal term
  - Velocity (momentum) is a state variable and is tracked in time

- 1D Wave equation

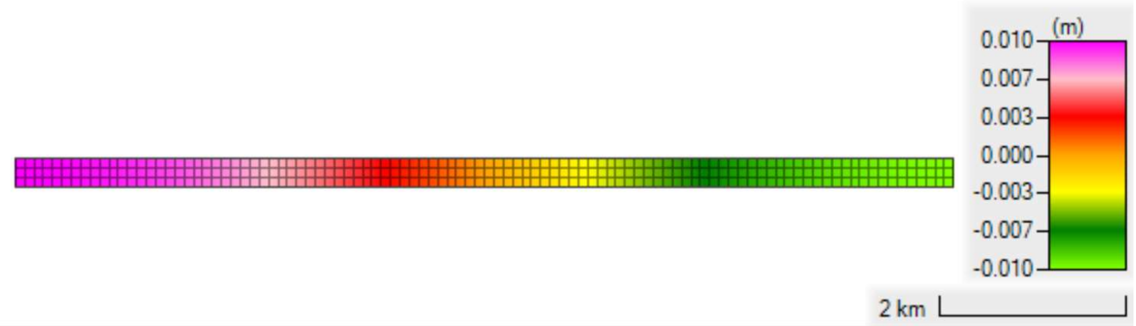
$$\frac{\partial z_s}{\partial t} = -h\frac{\partial u}{\partial x}$$

$$\frac{\partial u}{\partial t} = -g\frac{\partial z_s}{\partial x}$$

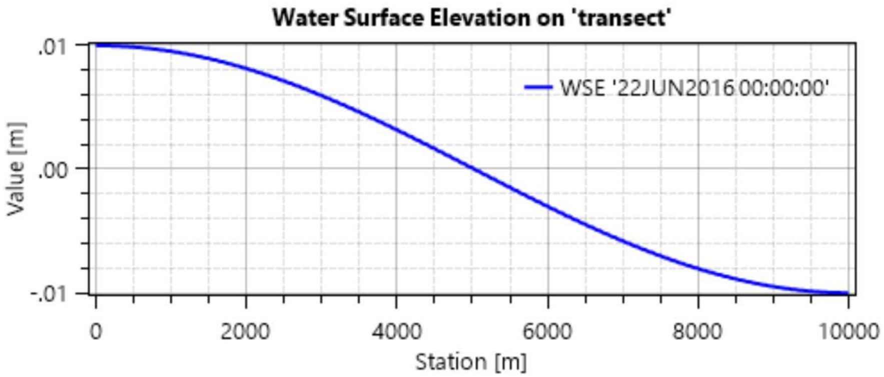
$$\Rightarrow \quad \frac{\partial^2 z_s}{\partial t^2} = c^2 \frac{\partial^2 z_s}{\partial x^2}$$

Example: Sloshing in a Rectangular Basin

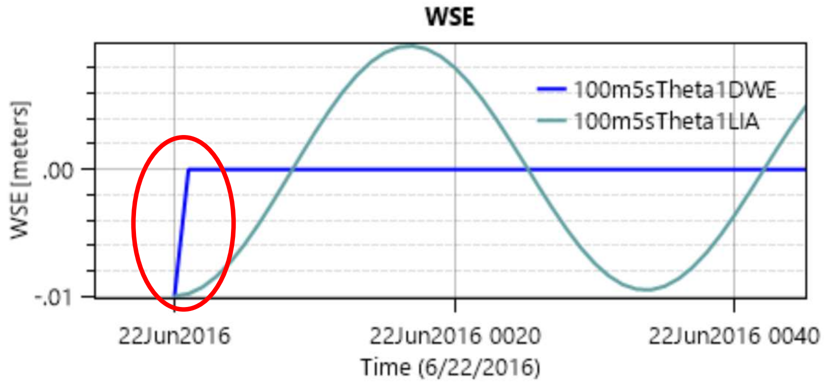
- Grid



- Initial Water Surface



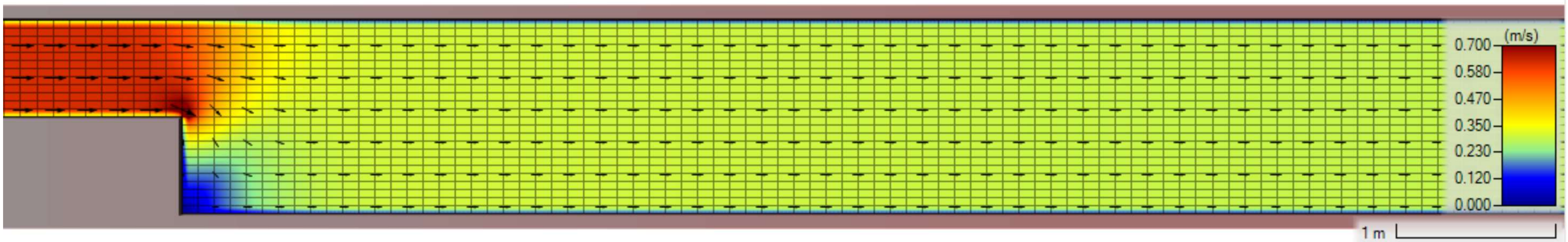
- Setup
  - **Diffusive-Wave**
  - **Local Inertial Approximation**
  - Implicit weighting factor: 1
  - Grid resolution: 100 m
  - Time step size: 5 s
- Time-series at one end



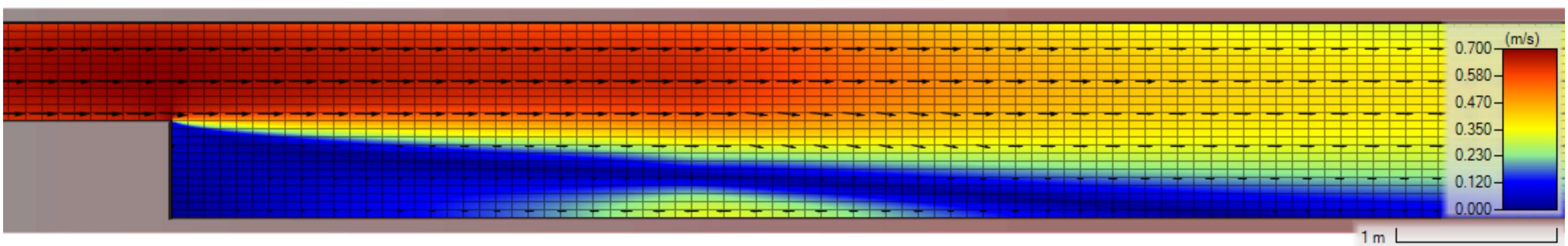
Advection

- Flume Experiment

DWE

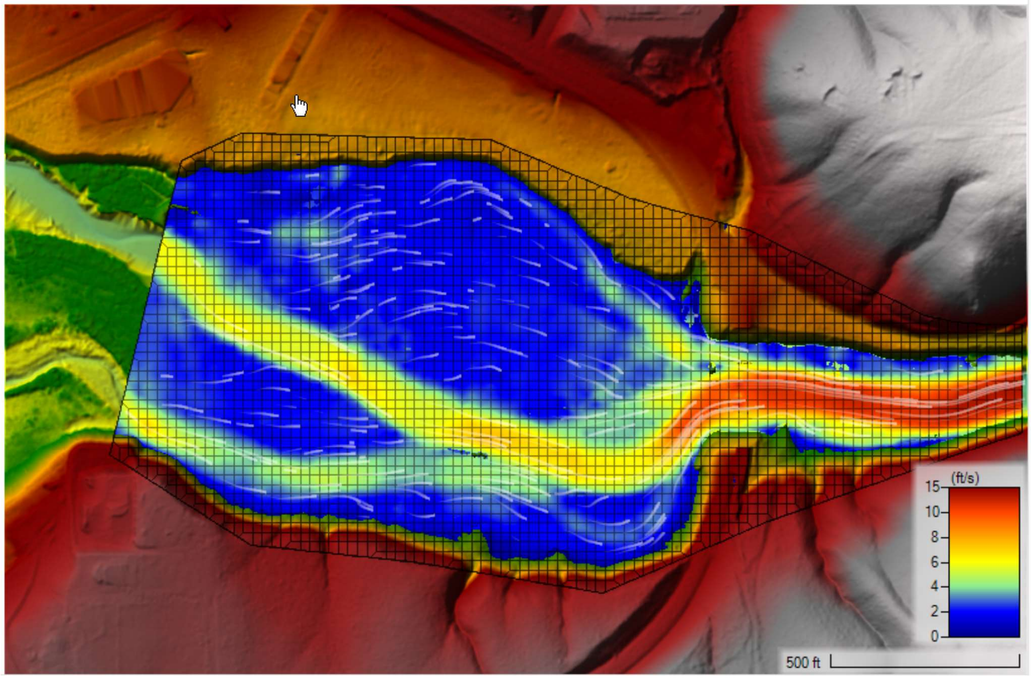


SWE

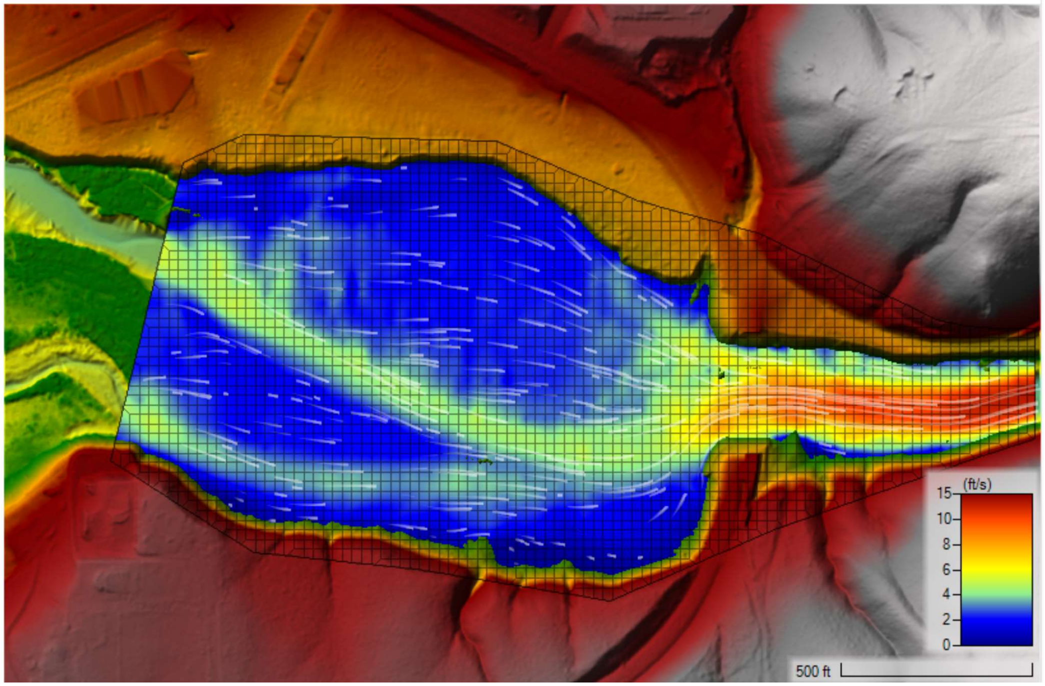


Advection

DWE

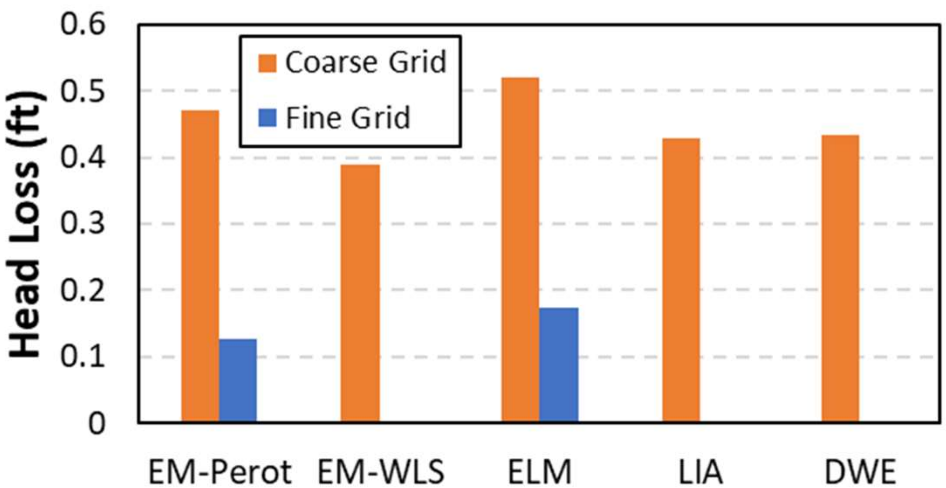


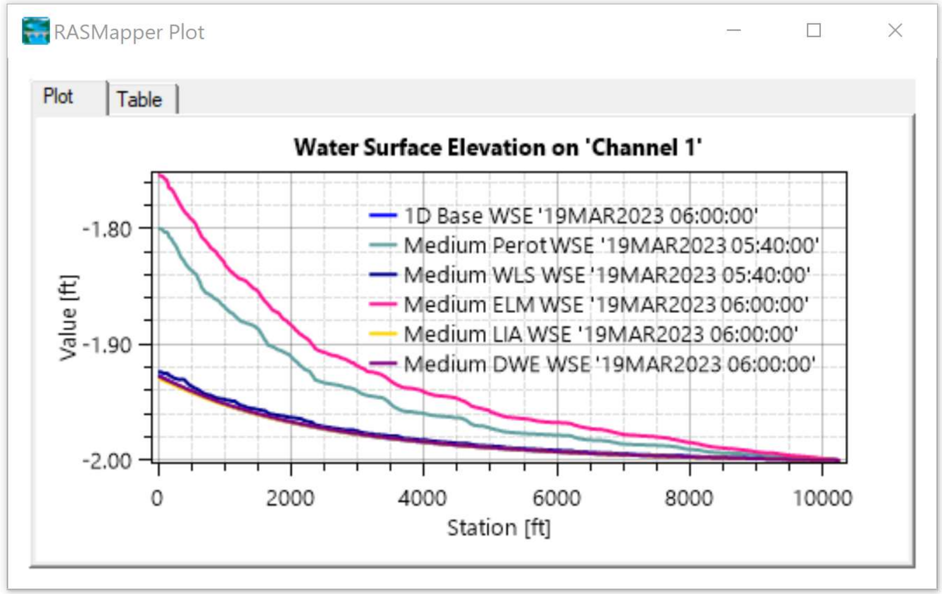
SWE-ELM

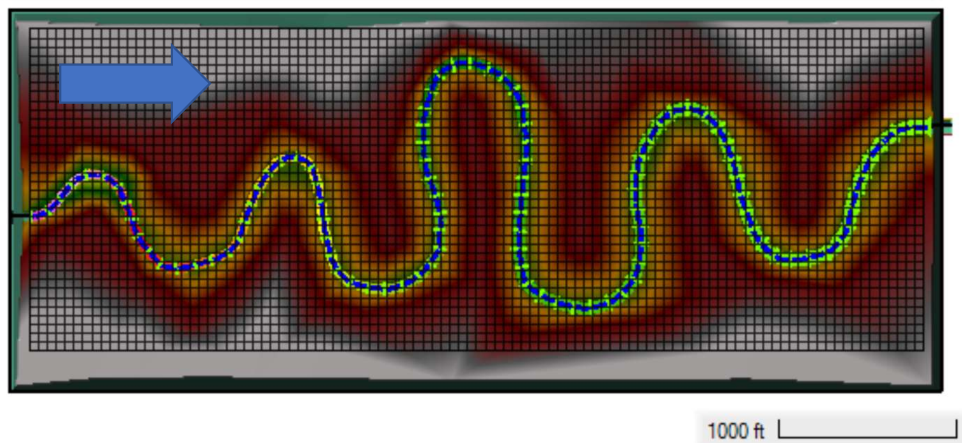


Numerical Dissipation

- Steady-flow in meandering river
- Computational errors especially from the advection term can produce artificial numerical dissipation
- Extremely coarse models benefit from ignoring the advection term







Diffusion of Momentum

- Non-conservative Formulation
  - Only option in Version 5.0.7 and earlier,
  - Optional in Version 6.0

$$\frac{D\boldsymbol{V}}{Dt} = -g\nabla z_s + \boxed{\nu_t \Delta \boldsymbol{V}} - \frac{\boldsymbol{\tau}_b}{\rho R}$$

- Conservative Formulation
  - Default in Version 6.0
  - Only option for Eulerian SWE solver

$$\frac{D\boldsymbol{V}}{Dt} = -g\nabla z_s + \boxed{\frac{1}{h}\nabla \cdot \left(\boldsymbol{\nu}_t h \nabla \boldsymbol{V}\right)} - \frac{\boldsymbol{\tau}_b}{\rho R}$$

$\Delta = \nabla^2$: Laplacian

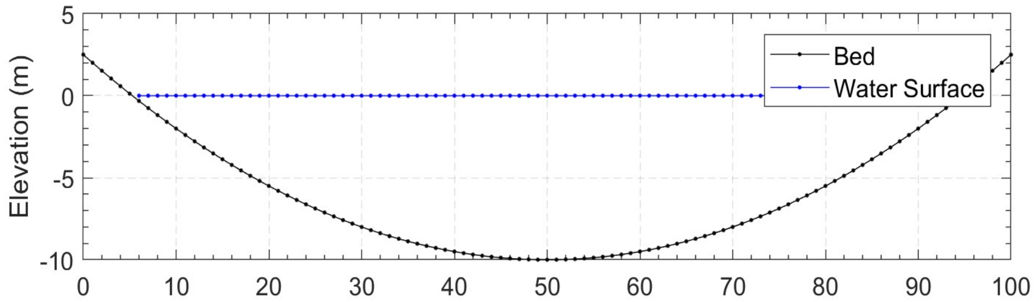
u_N : Face-normal velocity

ν_t : Turbulent eddy viscosity

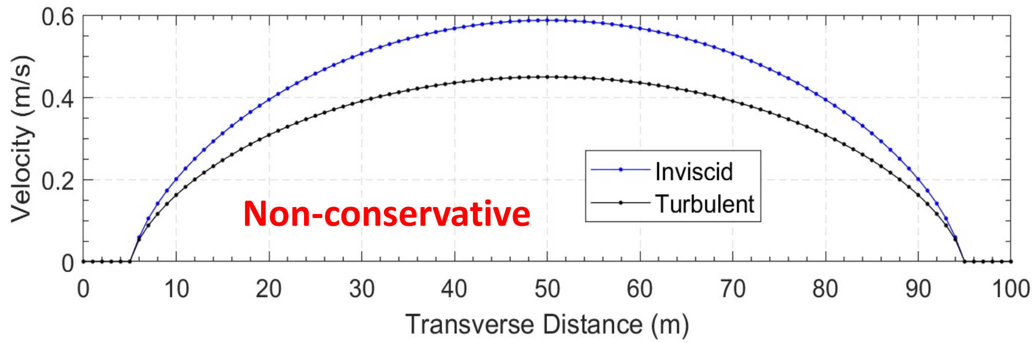
h : Water depth

c_f : Non-linear friction coefficient

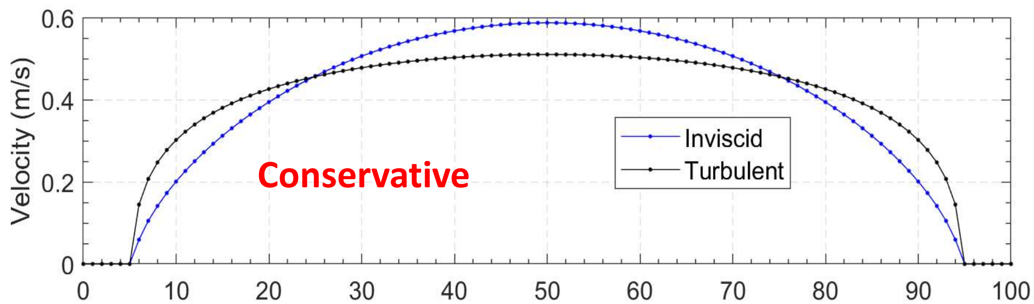
Mixing Term Formulation Comparison



Bathymetry and water level



Produces a net dissipation



Decreases velocities in middle of channel but increases velocities near banks

Eddy Viscosity: Turbulence Model

- Old: Parabolic $\nu_t = Du_* h$
  - Versions 5.0.7 and earlier
  - Isotropic (same in all directions)
  - 1 parameter: mixing coefficient D
- New: Parabolic-Smagorisnky

$$\boldsymbol{\nu}_t = \boldsymbol{D} u_* h + (C_s \Delta)^2 |\bar{S}|$$

u_* : Shear velocity

h : Water depth

D : Mixing coefficient

D_L : Longitudinal mixing coefficient

D_T : Transverse mixing coefficient

C_s : Smagorinsky coefficient

$$|\bar{S}| = \sqrt{2\left(\frac{\partial u}{\partial x}\right)^2 + 2\left(\frac{\partial v}{\partial y}\right)^2 + \left(\frac{\partial u}{\partial y} + \frac{\partial v}{\partial x}\right)^2} \qquad \boldsymbol{D} = \begin{bmatrix} D_{xx} & 0 \\ 0 & D_{yy} \end{bmatrix}$$

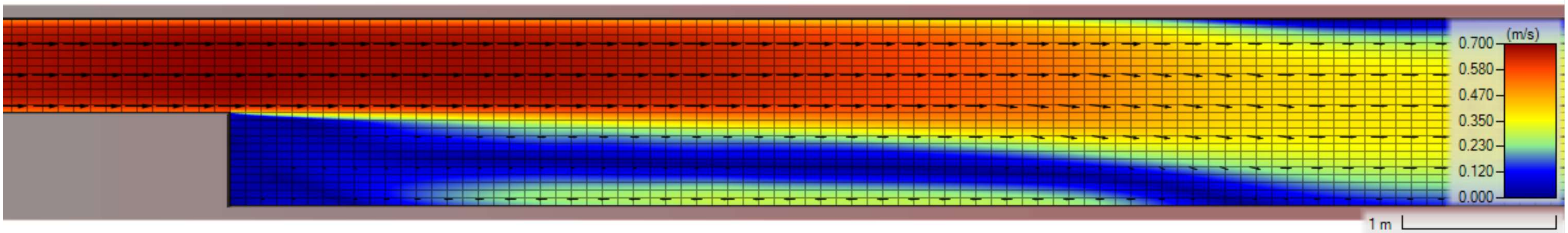
$$D_{xx} = D_L \cos^2\theta + D_{,T} \sin^2\theta$$

$$D_{yy} = D_L \sin^2\theta + D_T \cos^2\theta$$

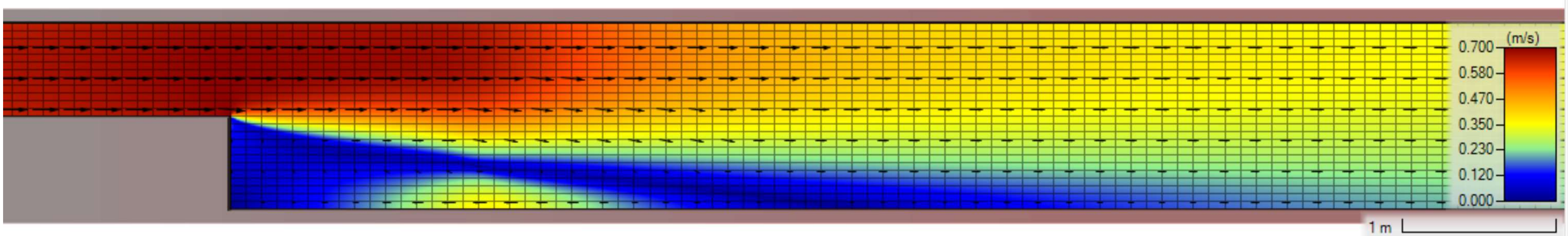
  - Default method in Version 6.0
  - Non-Isotropic (not the same in all directions)
  - 3 parameters: D_L, D_T, and C_s

Turbulence

- Low turbulence

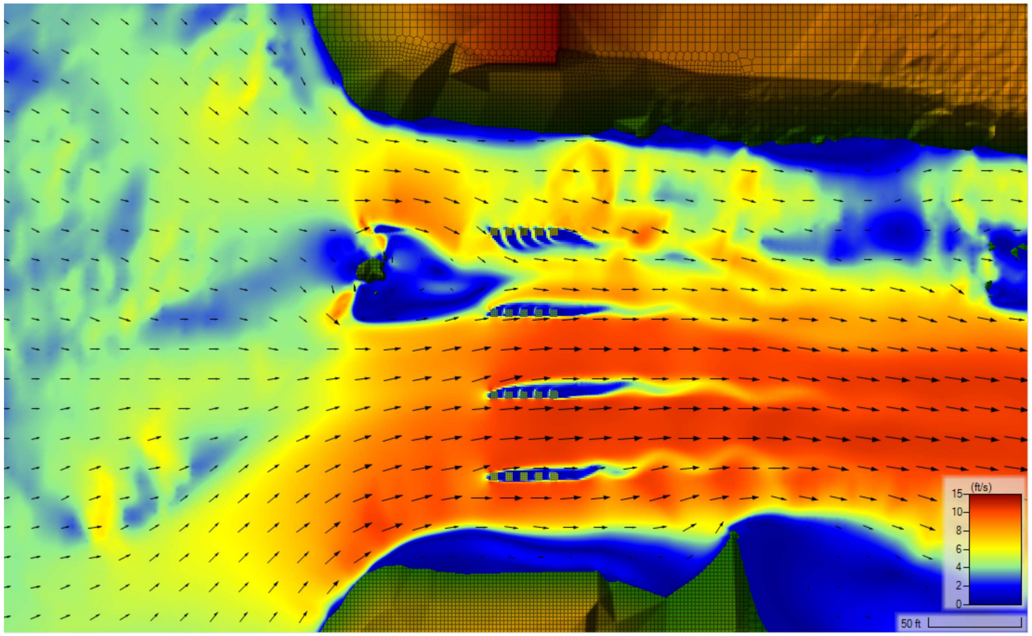


- High turbulence

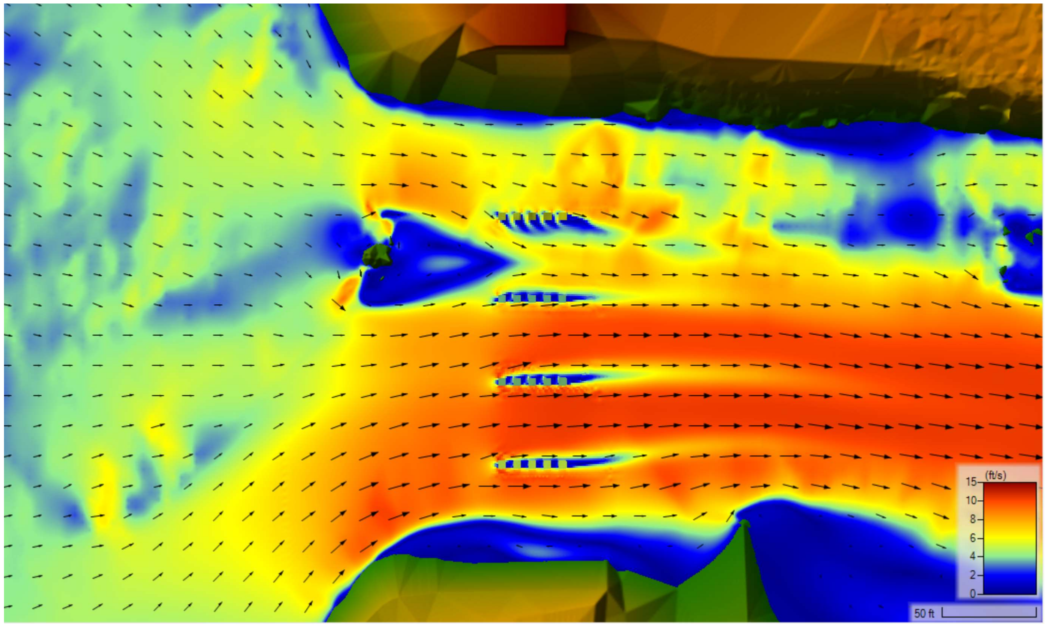


Turbulence

- No turbulence

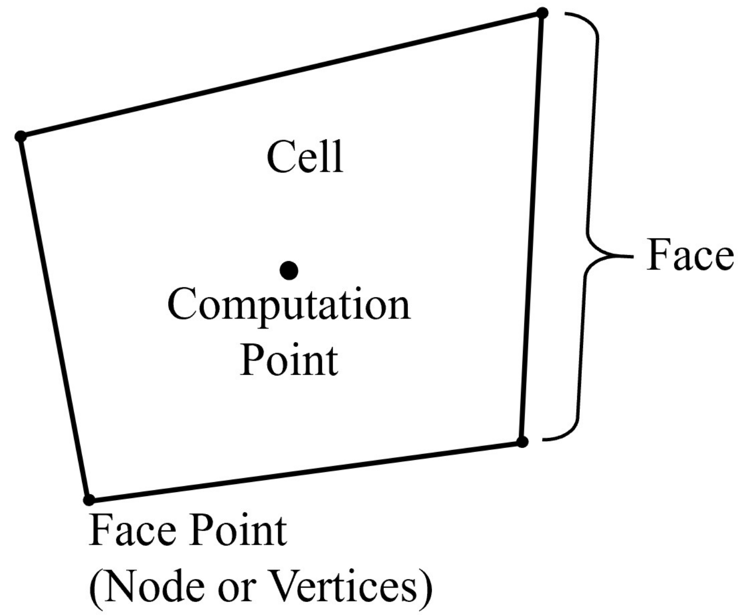


- With turbulence



Computational Mesh

- Mesh/grid can be unstructured
- Polygonal cells of up to 8 sides
- Cells must be concave
- Multiple 2D mesh can be run together or independently
- Grid Notation
  - Cells, Faces, Face Points (i.e. nodes or vertices), Computational Points, etc.
- State Variables
  - Cell Water levels
  - Face-normal Velocities



Numerical Methods

- Both DWE and SWE solvers are **Semi-implicit**
- Terms treated as:
  - Explicit: acceleration and diffusion terms
  - Semi-implicit: friction, flow divergence terms, and water level gradient
  - Fully-Implicit: pressure gradient term (for $\theta = 1$)
- By treating the "fast" pressure gradient term implicitly, the time step limitation based on the wave celerity can be removed
- Both DWE and SWE use **Finite-Difference** and **Finite-Volume** Methods
- Time integration: **Finite-Difference**
- Continuity Equation: **Finite-Volume**
- Momentum Equation: **Finite-Difference** (no control volume)

Implicit vs Explicit Time Stepping

- Explicit
  - Next state computed based solely on previous state
  - Easier to program and solve
  - Smaller time steps
  - Less robust
- Implicit
  - Next state computed based on previous state and next state
  - Harder to program and solve
  - Larger time steps
  - More robust

- Example:

$$\frac{\partial y}{\partial t} = F\left(y, t\right)$$

  - Explicit

$$y^{n+1} = y^n + \Delta t F^n$$

  - Implicit

$$y^{n+1} = y^n + \Delta t F^{n+1}$$

  - Semi-implicit

$$y^{n+1} = y^n + \Delta t\left[\theta F^{n+1} + (1-\theta) F^n\right]$$

Eulerian-Lagrangian vs. Eulerian SWE Solvers

- ELM-SWE
  - Only solver available in V5.0.7 and earlier
  - Default in V6.0
  - Not limited by Courant condition
  - Excellent stability
  - Can have momentum conservation problems around shocks or where the flow changes rapidly
- EM-SWE
  - New to V6.0 as an option
  - Limited to Courant less than 1.0
  - Good Stability
  - Improved momentum conservation for all flow conditions

Strength/Feature/Capability	**SWE-ELM**	**SWE-EM**
Larger Time Step	X	
Best Stability	X	
Courant Stability Criteria		X
Diffusion Stability Criteria		X
Computational Speed	X	
Wet/dry > 1 cell per time step	X	
Best Momentum Conservation		X
Non-Conservative Mixing	X	
Conservative Mixing	X	X
Wind	X	X

Solution Procedure

- System of equations

$$\boldsymbol{\Omega} + \boldsymbol{\Psi}\boldsymbol{Z} = \boldsymbol{b}$$

- Algorithm
  1. Compute Right-Hand-Side $\boldsymbol{b}$
     - Contains explicit terms: advection, diffusion, wind, etc.
  2. Outer Loop (Assembly and Updates)
     - Update linearized terms and variables including coefficient matrix $\boldsymbol{\Psi}$
  3. Inner Loop (Newton Iterations)

$$\boldsymbol{Z}^{m+1} = \boldsymbol{Z}^{m} - \left[\boldsymbol{\Psi} + \boldsymbol{A}^{m}\right]^{-1}\left(\boldsymbol{\Omega}^{m} + \boldsymbol{\Psi}\boldsymbol{Z}^{m} - \boldsymbol{b}\right)$$

$\boldsymbol{Z}$: Water level

$\boldsymbol{\Omega}$: Water volume

$\boldsymbol{\psi}$: Coefficient matrix

$\boldsymbol{b}$: Right-hand-side

m : Iteration index

$\boldsymbol{A}$: Diagonal matrix of cell wet surface areas

Boundary Conditions

- **Stage Hydrograph**. Upstream or downstream
- **Flow Hydrograph**. Upstream or downstream. Local conveyance and velocities computed automatically.
- **Normal Depth BC.** At downstream boundaries.
- **Rating Curve BC.**
- **Wind.** Only for shallow-water equations.
- **Precipitation, evapotranspiration, and infiltration.** Included as sources and sinks in the continuity equation.
- 1D reaches and 2D areas can be connected
- Multiple 2D areas can be connected to each other
- 2D areas can be connected to 1D lateral structures such as levees to simulate levee breaches

Computational Implementation

- Multiple 2D areas can be computed independently and simultaneously
- All solvers are can be run on multiple cores
- 2D solvers and parameters can be selected independently for each 2D area
- A partial grid solution keeps track of active portion of mesh and only computes the solution for active portion significantly reducing computational times.

Thank You!

HEC-RAS Website:
https://www.hec.usace.army.mil/software/hec-ras/

Online Documentation:
https://www.hec.usace.army.mil/confluence/rasdocs

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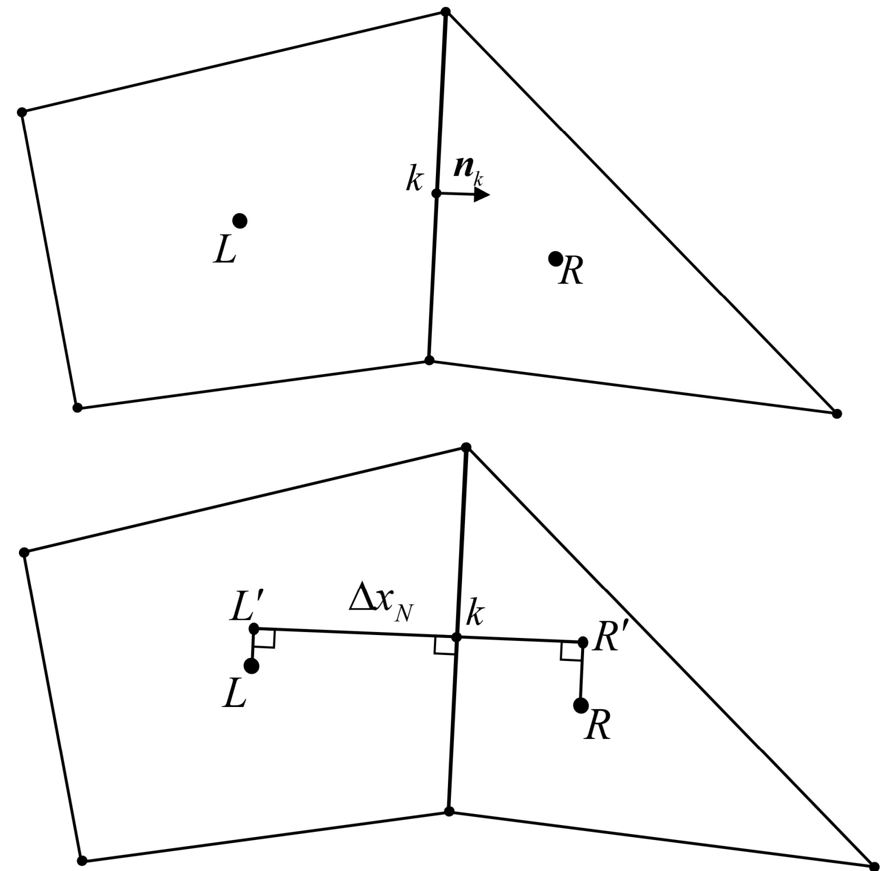
Face Water Surface Gradient

- Face-Normal Gradient

$$\nabla z_s \cdot \boldsymbol{n}_k = \frac{\partial z_s}{\partial N} \approx \frac{z_{s,R} - z_{s,L}}{\Delta x_N}$$

  - Uses **Cell Centroids** and **NOT** the **Computation Points**
  - Future versions may include non-orthogonal
  - Compact two-point stencil is computationally efficient and robust
  - Important to have a good quality mesh to reduce errors



Momentum Conservation

- Momentum conservation is directionally invariant
- Only “face-normal” component is needed at faces so

$$\frac{\partial u_N}{\partial t} + (\boldsymbol{V} \cdot \nabla) u_N - f_c u_T = -g \frac{\partial z_s}{\partial N} + \frac{1}{h} \nabla \cdot (\boldsymbol{v}_t h \nabla u_N) - \frac{\tau_{b,N}}{\rho R} + \frac{\tau_{s,N}}{\rho h}$$

where u_N is the velocity in the N direction

Face-Tangential Velocity

- Tangential velocities are computed on left and right of face with a Least-squares Formulation

$$S_R = \sum_{k \in R}^{3} \left(\boldsymbol{V}_R \cdot \boldsymbol{n}_k - (u_N)_k \right)^2 \qquad S_L = \sum_{k \in L}^{3} \left(\boldsymbol{V}_L \cdot \boldsymbol{n}_k - (u_N)_k \right)^2$$

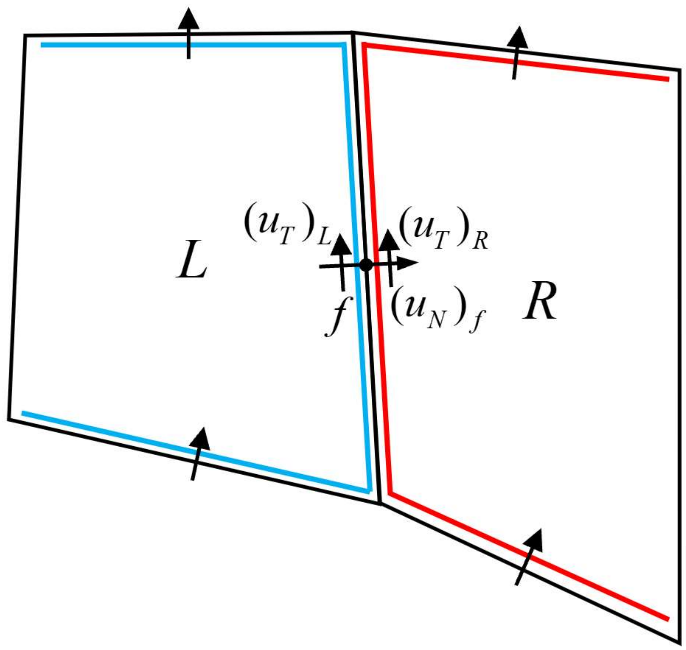
- Of the left and right reconstructed velocities, only the tangential component is used, because the normal component is known

$$(u_T)_R = \boldsymbol{V}_R \cdot \boldsymbol{t}_f \qquad (u_T)_L = \boldsymbol{V}_L \cdot \boldsymbol{t}_f$$

- Average face-tangential velocity computed as

$$(u_T)_f = \frac{(u_T)_R + (u_T)_L}{2}$$

Double-C Stencil



Discretization

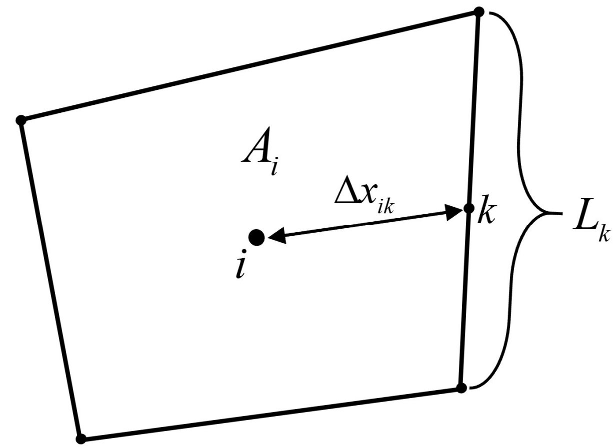
- Cell Velocity Gradient (x-direction)
  - Gauss' Divergence Theorem

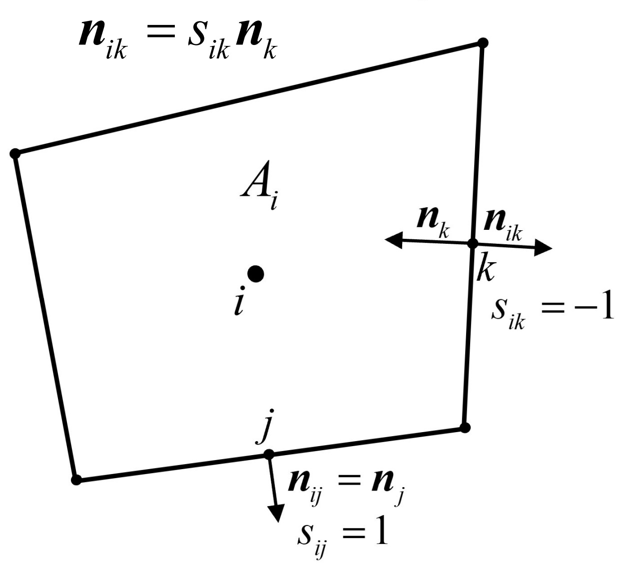
$$\nabla u_i = \frac{1}{A_i}\int_A \nabla u \, \mathrm{d}A = \frac{1}{A_i}\oint_L u\boldsymbol{n} \, \mathrm{d}L = \frac{1}{A_i}\sum_{k\in i} u_k \boldsymbol{n}_{ik} L_k$$

  - Needed tor turbulence modeling
- Cell Velocity
  - Perot's Method

$$\boldsymbol{V}_i = \frac{1}{A_i}\sum_{k\in i} \Delta x_{ik} L_k \boldsymbol{n}_k (u_N)_k$$

  - Needed for the conservative form of the mixing term and for Eulerian advection





Discretization: Laplacian

- Node Laplacian

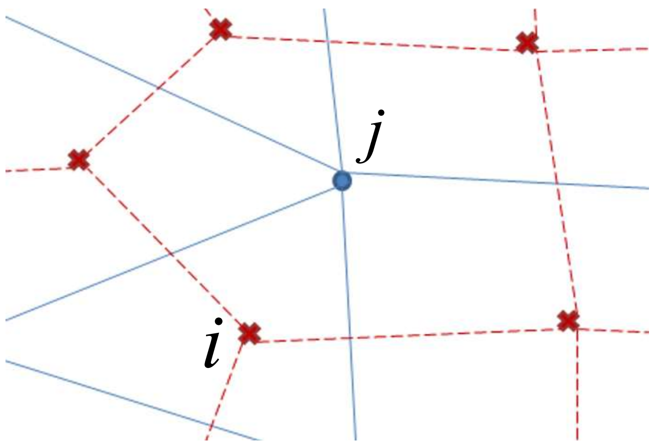
$$\left(\nabla^2 \boldsymbol{V}\right)_j = \left[\nabla \cdot \left(\nabla \boldsymbol{V}\right)\right]_j \approx \sum_i d_i \left(\nabla \boldsymbol{V}\right)_i$$

$$\left(\nabla \boldsymbol{V}\right)_i = \sum_k c_k \boldsymbol{V}_k$$

i : Cells

j : Nodes

k : Faces



- Used only by non-conservative turbulence

Backtracking

$$\frac{\partial u_N}{\partial t} + \left(\boldsymbol{V} \cdot \nabla\right) u_N = \frac{D u_N}{Dt} \approx \frac{u_N^{n+1} - u_X^n}{\Delta t}$$

$$\frac{d\boldsymbol{x}_P}{dt} = \boldsymbol{V}(\boldsymbol{x}, t)$$

1. Interpolate node velocities from faces
2. Set starting location and remaining time as f and $T_R = \Delta t$
3. From starting location and velocity, find location B
4. Compute time to location B: $T_B = (\boldsymbol{x}_A - \boldsymbol{x}_B)\boldsymbol{V}_A^{-1}$
5. Interpolate velocity at location B: $\boldsymbol{V}_B = w_{n1}\boldsymbol{V}_{n1} + w_{n2}\boldsymbol{V}_{n2}$

if $T_B > T_R$

6. Set $A = B$, $T_R = T_R - T_B$, and go to step 3

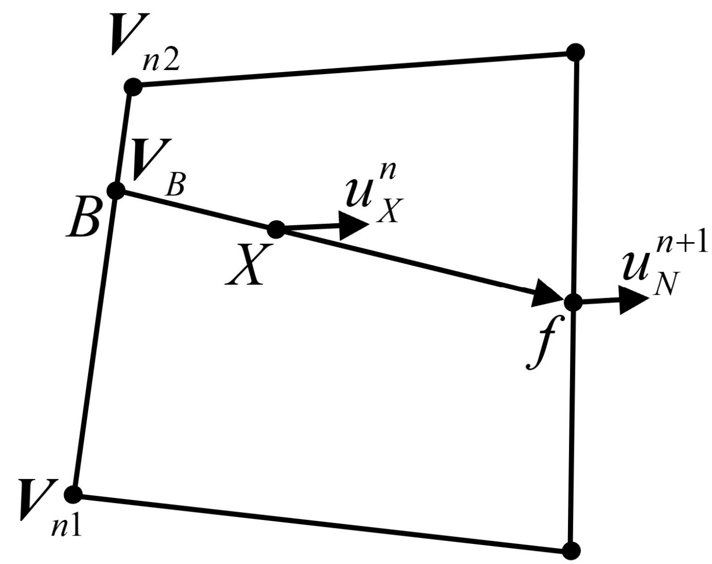
else

7. Find location X as $\boldsymbol{x}_X = \boldsymbol{x}_f - T_R \boldsymbol{V}_A$
8. Interpolate velocity vector at X

$$\boldsymbol{V}_X = T_B^{-1}\left[T_R \boldsymbol{V}_B + \left(T_B - T_R\right)\boldsymbol{V}_A\right]$$

9. Compute advective velocity

$$u_X = \boldsymbol{n}_f \cdot \boldsymbol{V}_X$$



Fractional Step Method (ELM only)

- Coriolis Term approximated as

$$f_c \boldsymbol{k} \times \boldsymbol{V} \approx \begin{pmatrix} f\left[(1-\theta) f v_X^n + \theta v^{n+1}\right] \\ -f\left[(1-\theta) f u_X^n + \theta u^{n+1}\right] \end{pmatrix}$$

where

f : Coriolis Parameter

θ : Implicit weighting factor

$\boldsymbol{k}$: Unit vector in the vertical direction

$\boldsymbol{V} = (u, v)^T$: Velocity at face

$\boldsymbol{V}_X = (u_X, v_X)^T$: Velocity at face at location X

- First (Coriolis) Step

$$\begin{pmatrix} 1 & \theta \Delta t f \\ \theta \Delta t f & 1 \end{pmatrix} \begin{pmatrix} u^* \\ v^* \end{pmatrix} = \begin{pmatrix} u_X^n + (1-\theta)\Delta t f v_X^n \\ v_X^n + (1-\theta)\Delta t f u_X^n \end{pmatrix} \qquad \boldsymbol{V}^* = \begin{pmatrix} u^* \\ v^* \end{pmatrix}$$

- Second Step includes all other terms

Eulerian-Lagrangian Momentum Equation

- Semi-discrete form (2nd Fractional Step)

$$\frac{u_N^{n+1}-u_N^{*}}{\Delta t}=-g\frac{\partial z_s^{n+\theta}}{\partial N}+\left[\frac{1}{h}\nabla\cdot\left(\nu_t h^n \nabla u_N\right)\right]_X^n - c_f u_N^{n+1}+\frac{\tau_{s,N}}{\rho h_f^n}$$

where

$$z_a^{n+\theta}=(1-\theta)z_s^n+\theta z_s^{n+1}$$

$$u_N^{*}=\boldsymbol{V}^{*}\cdot\boldsymbol{n}_f$$

- Velocity $\boldsymbol{V}^{*}$ includes Coriolis
- Mixing term is interpolated at backtracking location X and based on previous time step velocity field
- Friction term is semi-implicit

Eulerian Momentum Equation

- Semi-discrete form

$$\frac{u_N^{n+1} - u_N^n}{\Delta t} + \left(\boldsymbol{V}^n \cdot \nabla\right) u_N^n - f u_T^n = -g \frac{\partial z_s^{n+\theta}}{\partial N} + \left[\frac{1}{h} \nabla \cdot \left(\nu_t h \nabla u_N\right)\right]_f^n - c_f u_N^{n+1} + \frac{\tau_{s,N}}{\rho h_f^n}$$

where

$$z_s^{n+\theta} = (1-\theta) z_s^n + \theta z_s^{n+1} \qquad \overline{h}_f = \alpha_f^L h_L + \alpha_f^R h_R$$

- Coriolis term computed at face f and is explicit
- No fractional step method like ELM solver
- Mixing term is computed at face f and is explicit
- Friction and pressure gradient terms are semi-implicit

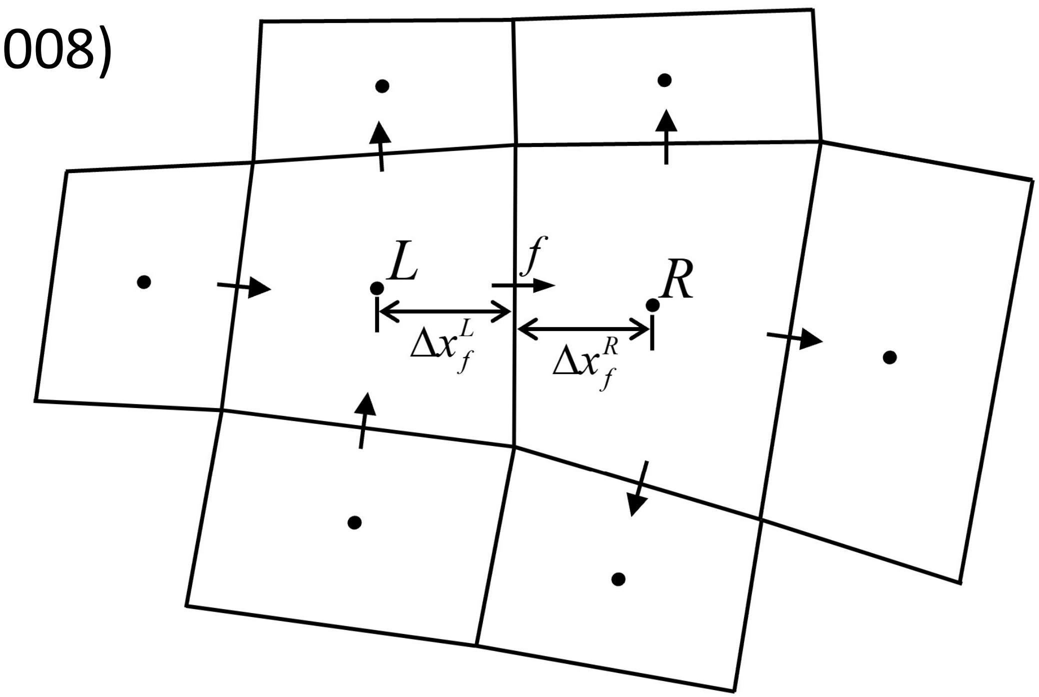
Discretization: Eulerian Advection

- Approach from Kramer and Stelling (2008)

$$(\boldsymbol{V}\cdot\nabla)u_N \approx \frac{\alpha_f^L}{\bar{h}_f A_L}\sum_{k\in L} s_{Lk}Q_k\left[\boldsymbol{V}_k^u\cdot\boldsymbol{n}_f-(u_N)_f\right]$$
$$+\frac{\alpha_f^R}{\bar{h}_f A_R}\sum_{k\in R} s_{Rk}Q_k\left[\boldsymbol{V}_k^u\cdot\boldsymbol{n}_f-(u_N)_f\right]$$

$$\alpha_f^L=\frac{\Delta x_f^L}{\Delta x_f^L+\Delta x_f^R} \qquad \bar{h}_f=\alpha_f^L h_L+\alpha_f^R h_R$$

$$\alpha_f^R=1-\alpha_f^L$$



- Courant-Freidrichs-Lewy (CFL) Condition

$$C=\frac{U\Delta t}{\Delta x}\le 1$$

Discretization: Mixing Term

- Non-Conservative Form

$$\boldsymbol{v}_t \nabla^2 u_N \Big|_f \approx v_{t,f}^n \left(\nabla^2 \boldsymbol{V}\right)_X^n \cdot \boldsymbol{n}_f$$

- Conservative Form

$$\frac{1}{h}\nabla\cdot\left(\boldsymbol{v}_t h \nabla u_N\right)\Big|_f \approx \frac{\alpha_f^L}{\bar{h}_f A_L}\sum_{k\in L} A_k v_{t,k} \frac{\boldsymbol{n}_f\cdot\left(\boldsymbol{V}_j - \boldsymbol{V}_L\right)}{\Delta x_{L,j}} + \frac{\alpha_f^R}{\bar{h}_f A_R}\sum_{k\in R} A_k v_{t,k} \frac{\boldsymbol{n}_f\cdot\left(\boldsymbol{V}_j - \boldsymbol{V}_R\right)}{\Delta x_{R,j}}$$

- Discretization same for both ELM and EM solvers
- Approximate Stability Criteria for EM solver

$$\frac{v_t \Delta t}{\Delta x^2} \leq \frac{1}{2}$$

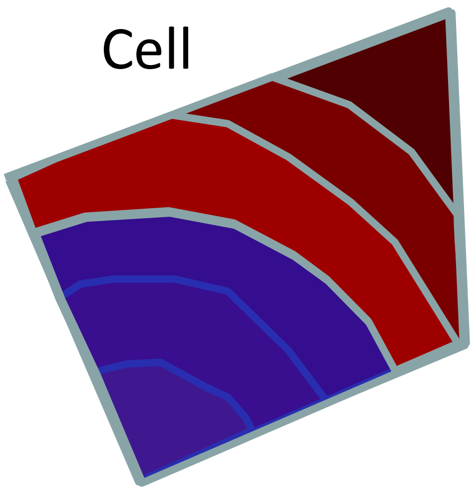
- ELM interpolates term to location X

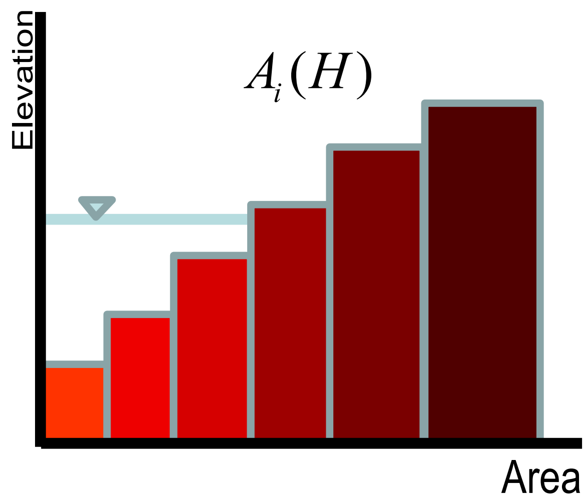
$$\left[\frac{1}{h}\nabla\cdot\left(\boldsymbol{v}_t h \nabla u_N\right)\right]_X^n$$

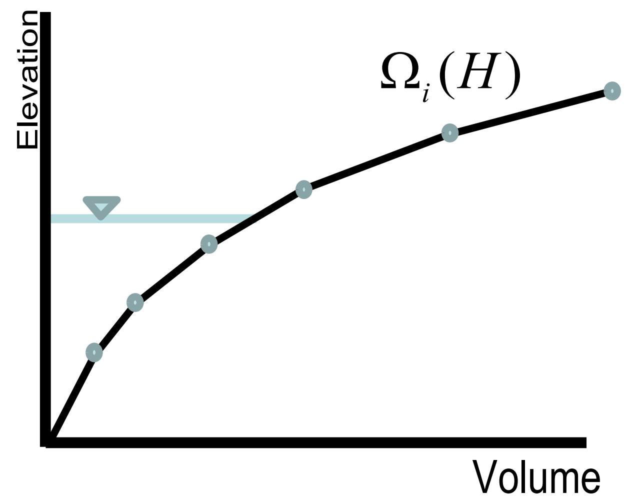
Subgrid Modeling

- Problem
  - Water levels usually vary much more smoothly than the terrain
  - Unfeasible to resolve every detail of the terrain with the computational mesh
- Approach
  - Utilize a grid resolution sufficient to resolve the hydraulics
  - Capture the details of the subgrid terrain through hydraulic properties tables

Subgrid Bathymetry: Cells

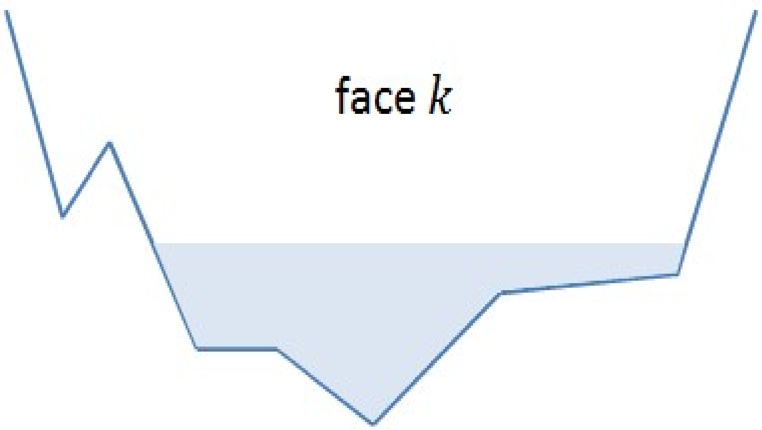


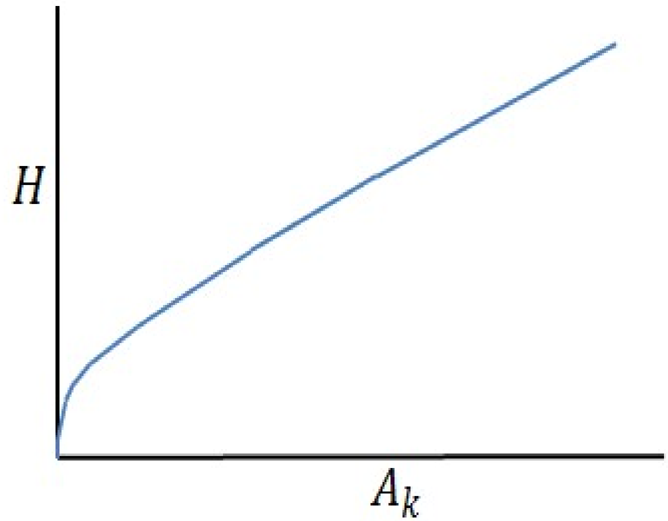




Subgrid Bathymetry: Faces

- Faces treated similar to cells
- Hydraulic property tables computed
  - Wetted length
  - Wetted Perimeter
  - Area





Benefits of Subgrid Bathymetry