

# Introduction to 2D Hydraulics Equations

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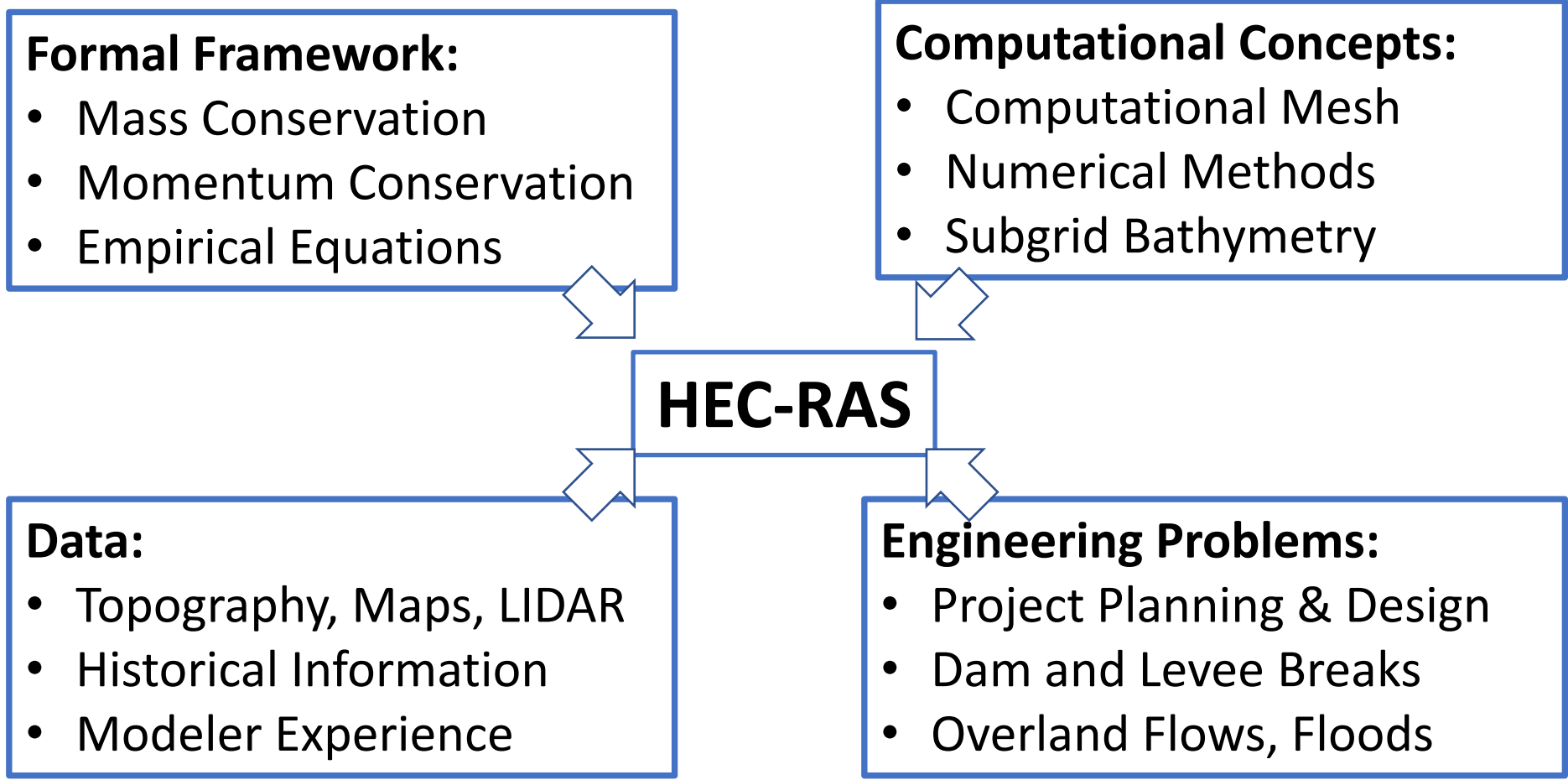
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# Hydraulic Modeling





# Outline

- Mass Conservation (Continuity)
- Momentum Conservation (Depth-Averaged)
  - Acceleration
  - Coriolis term
  - Hydrostatic pressure
  - Turbulent mixing
  - Friction
- Diffusion Wave Equation
- Numerical Methods



# Mass Conservation

- Assuming a constant water density

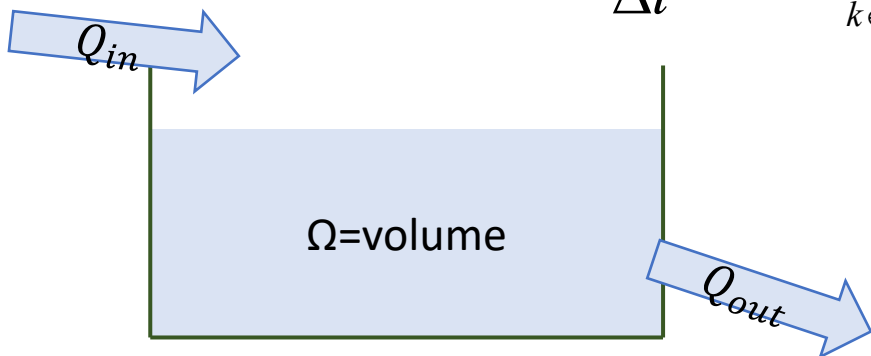
$$\frac{\partial H}{\partial t} + \nabla \cdot (h\mathbf{V}) = q$$

- Integrating over a computational cell

$$\frac{\partial}{\partial t} \iiint_{\Omega} d\Omega + \iint_S (\mathbf{V} \cdot \mathbf{n}) dS = Q$$

- Finite-Volume Discretization

$$\frac{\Omega_i^{n+1} - \Omega_i^n}{\Delta t} + \sum_{k \in i} (\mathbf{V}_k \cdot \mathbf{n}_{ik}) A_k = Q$$



Change in volume in a system balances with flow through boundaries

$H$  : Water surface elevation

$h$  : Water depth

$q$  : Water source/sink

$\Omega_i$  : Cell water volume

$A_k$  : Face area

$\mathbf{V}_k$  : Face velocity

$\mathbf{n}_{ik}$  : Outward face-normal unit vector

$\Delta t$  : Time step



# Momentum Conservation

- Momentum Equation (non-conservative form)

$$\frac{\partial \mathbf{V}}{\partial t} + (\mathbf{V} \cdot \nabla) \mathbf{V} + f \mathbf{k} \times \mathbf{V} = -g \nabla H + \frac{1}{h} \nabla \cdot (\mathbf{v}_t h \nabla \mathbf{V}) - c_f \mathbf{V} + \frac{\boldsymbol{\tau}_s}{\rho h}$$

- From Newton's 2<sup>nd</sup> Law of motion
- Assumes constant water density, small vertical velocities, hydrostatic pressure, etc.
- Non-linear and a function of both velocity and water levels
- Continuity and Momentum Equations are the Shallow Water Equations or sometimes referred to as the "Full Momentum" equations in HEC-RAS



# Acceleration and Total Derivative

- **Eulerian:** Frame of reference is fixed in space and time

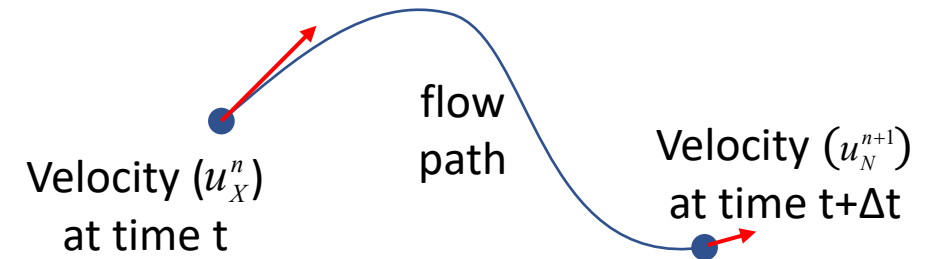
$$\frac{\partial \mathbf{V}}{\partial t} + (\mathbf{V} \cdot \nabla) \mathbf{V}$$

- Easier to compute
- Time-step restricted by Courant condition

- **Lagrangian:** Frame of reference moves with total derivative along flow path

$$\frac{\partial \mathbf{V}}{\partial t} + (\mathbf{V} \cdot \nabla) \mathbf{V} = \frac{D\mathbf{V}}{Dt} = \frac{\mathbf{V}^{n+1} - \mathbf{V}_X^n}{\Delta t}$$

- More expensive to compute
- Allows larger time-steps



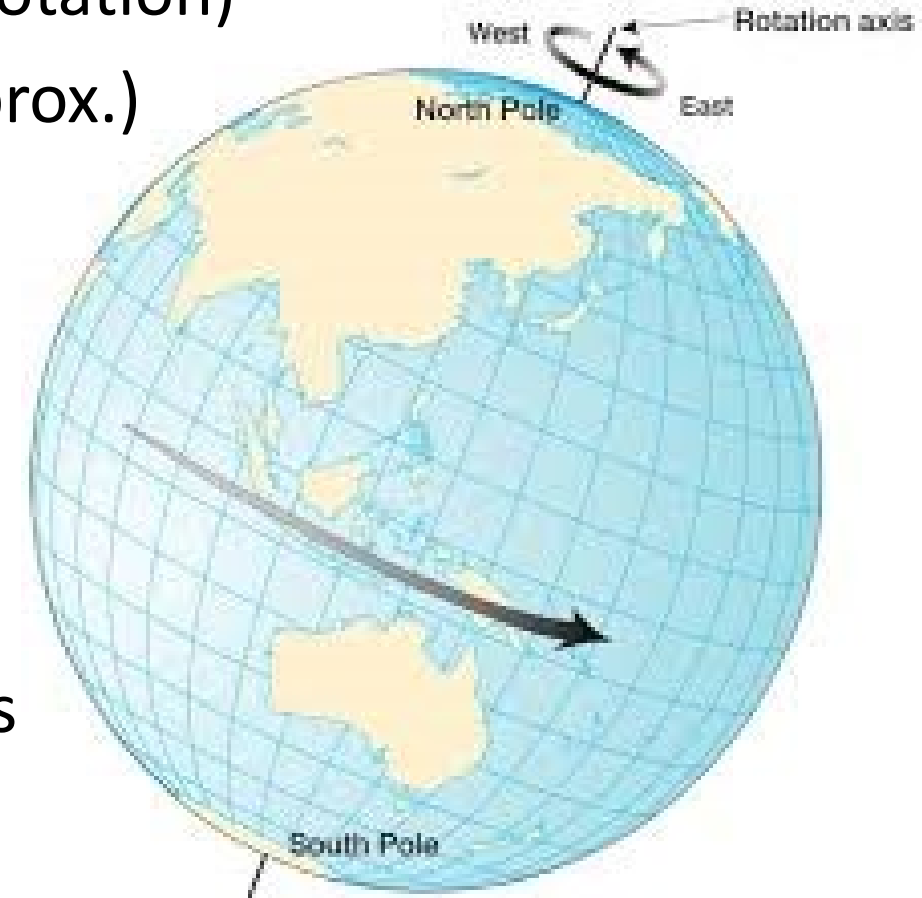


# Coriolis Acceleration

- Effect of rotating frame of reference (earth's rotation)
- Constant for the each 2D domain (f-plane approx.)

$$f = 2 \omega \sin \varphi$$

- $\omega$ : sidereal angular velocity of the Earth
- $\varphi$ : latitude. Positive for northern hemisphere. Negative for southern hemisphere
- Coriolis acceleration disabled by default to save computational time
- Negligible for most river and flood simulations
- When to enable Coriolis term?
  - Large domains
  - Higher latitudes





# Hydrostatic Pressure

- Assumes vertical water accelerations are small compared to gravity
- Total pressure is

$$P = P_{atm} + \rho g(H - z)$$

- $P_{atm}$ : atmospheric pressure (assumed to be constant)
  - $\rho$  : constant water density
  - $g$ : gravity acceleration constant
  - $H$ : water surface elevation
  - $z$ : vertical coordinate
- Pressure gradient

$$\frac{\partial P}{\partial x} = \rho g \frac{\partial H}{\partial x}$$





# Diffusion of Momentum

- Non-conservative Formulation
  - Only option in Version 5.0.7 and earlier,
  - Optional in Version 6.0

$$\frac{DV}{Dt} = -g\nabla H + \boxed{v_t \Delta V} - c_f V$$

- Conservative Formulation
  - Default in Version 6.0
  - Only option for Eulerian SWE solver

$$\frac{DV}{Dt} = -g\nabla H + \boxed{\frac{1}{h} \nabla \cdot (v_t h \nabla V)} - c_f V$$

$\Delta = \nabla^2$  : Laplacian

$u_N$  : Face-normal velocity

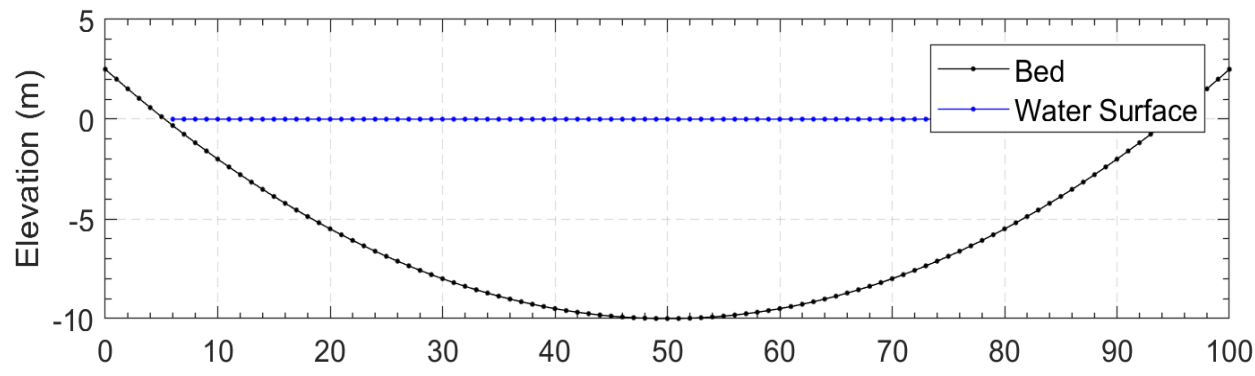
$v_t$  : Turbulent eddy viscosity

$h$  : Water depth

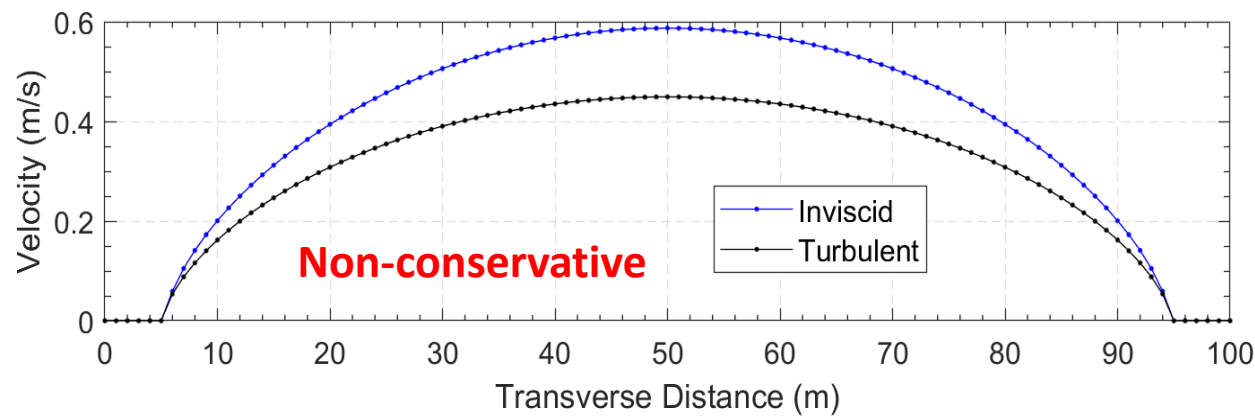
$c_f$  : Non-linear friction coefficient



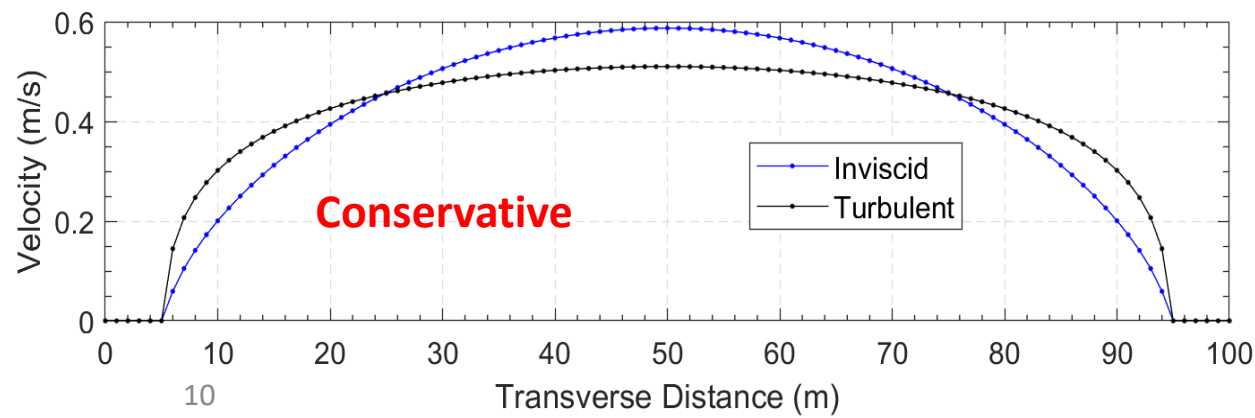
# Mixing Term Formulation Comparison



Bathymetry and water level



Produces a net dissipation



Decreases velocities in middle of channel but increases velocities near banks



# Eddy Viscosity: Turbulence Model

- Old: Parabolic  $\nu_t = Du_*h$ 
  - Versions 5.0.7 and earlier
  - Isotropic (same in all directions)
  - 1 parameter: mixing coefficient  $D$
- New: Parabolic-Smagorinsky

$u_*$  : Shear velocity

$h$  : Water depth

$D$  : Mixing coefficient

$D_L$  : Longitudinal mixing coefficient

$D_T$  : Transverse mixing coefficient

$C_s$  : Smagorinsky coefficient

$$\nu_t = \mathbf{D}u_*h + (C_s\Delta)^2 |\bar{S}|$$

$$|\bar{S}| = \sqrt{2\left(\frac{\partial u}{\partial x}\right)^2 + 2\left(\frac{\partial v}{\partial y}\right)^2 + \left(\frac{\partial u}{\partial y} + \frac{\partial v}{\partial x}\right)^2} \quad \mathbf{D} = \begin{bmatrix} D_{xx} & 0 \\ 0 & D_{yy} \end{bmatrix} \quad \begin{aligned} D_{xx} &= D_L \cos^2 \theta + D_T \sin^2 \theta \\ D_{yy} &= D_L \sin^2 \theta + D_T \cos^2 \theta \end{aligned}$$

- Default method in Version 6.0
- Non-Isotropic (not the same in all directions)
- 3 parameters:  $D_L$ ,  $D_T$ , and  $C_s$



# Bottom Friction

- Resisting force due to relative motion of fluid against the bed
- Non-linear friction coefficient

$$c_f = \frac{n^2 g |V|}{R^{4/3}}$$

- $n$  :Manning coefficient
- $g$ : gravity acceleration constant
- $|V|$  : velocity magnitude
- $R$ : hydraulic radius

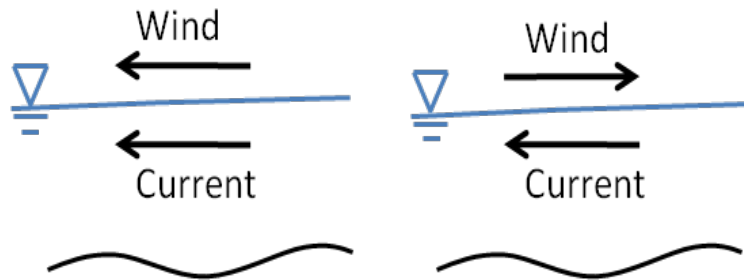
# Wind Stress

- Surface Stress is given by

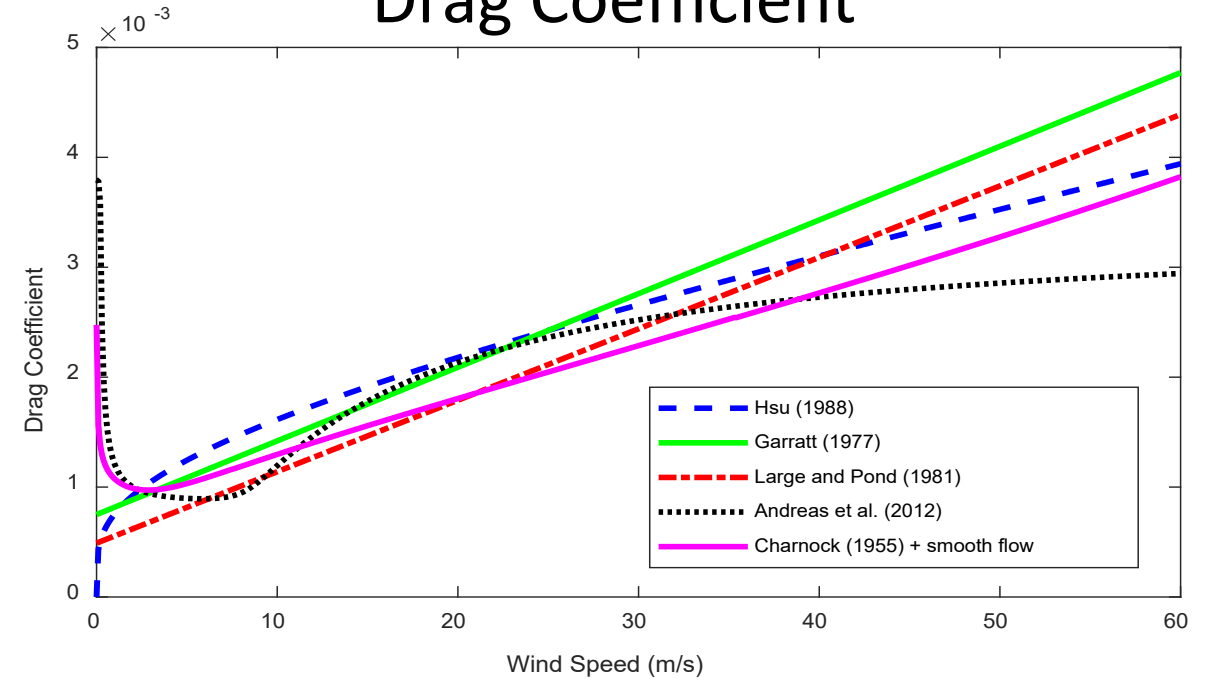
$$\tau_s = \rho_a C_D |\mathbf{W}_{10}| \mathbf{W}_{10}$$

- Wind Reference Frame

$$\mathbf{W}_{10} = \begin{cases} \mathbf{W}_{10}^E - \mathbf{V} & \text{for Lagrangian} \\ \mathbf{W}_{10}^E & \text{for Eulerian} \end{cases}$$



## Drag Coefficient





# Diffusion-Wave Approximation

- Ignoring the acceleration, Coriolis, mixing, and wind-forcing terms, the momentum equation reduces to

$$\frac{n^2}{R^{4/3}} |\mathbf{V}| \mathbf{V} = -\nabla H$$

- Dividing both sides by the square of its norm

$$\mathbf{V} = -\frac{R^{2/3}}{n} \frac{\nabla H}{|\nabla H|^{1/2}}$$

- Inserting the above equation into the Continuity Equation leads to the Diffusion-Wave Equation (DWE)

$$\frac{\partial H}{\partial t} = \nabla \cdot (\beta \nabla H) + q \quad \beta = \frac{R^{2/3} h}{n |\nabla H|^{1/2}}$$



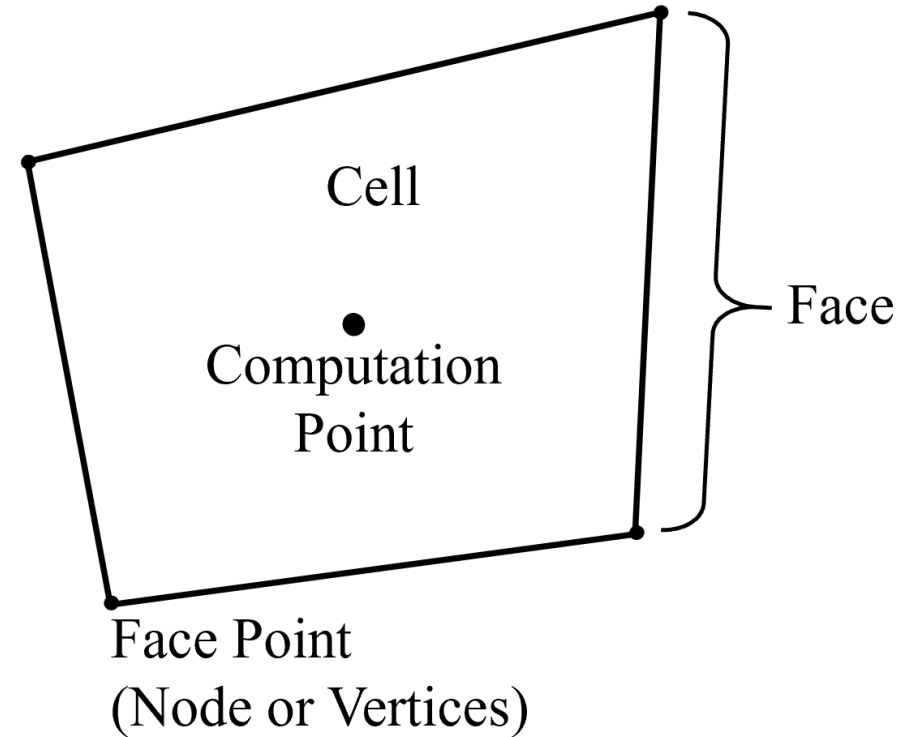
# SWE vs. DWE



- Use SWE for:
  - Flows with dynamic changes in acceleration
  - Studies with important wave effects, tidal flows
  - Detail solution of flows around obstacles, bridges or bends
  - Simulations influenced by Coriolis, mixing, or wind
  - To obtain high-resolution and detailed flows
- Use DWE for:
  - Flow is mainly driven by gravity and friction
  - Fluid acceleration is monotonic and smooth, no waves
  - To compute approximate global estimates such as flood extent
  - To assess approximate effects of dam breaks
  - To assess interior areas due to levee breaches
  - For quick estimations or preliminary runs

# Computational Mesh

- Mesh/grid can be unstructured
- Polygonal cells of up to 8 sides
- Cells must be concave
- Multiple 2D mesh can be run together or independently.
- Grid Notation
  - Cells, Faces, Face Points (i.e. nodes or vertices), Computational Points, etc.







# Numerical Methods

- Both DWE and SWE solvers are **Semi-implicit**
- Terms treated as:
  - Explicit: acceleration and diffusion terms
  - Semi-implicit: friction, flow divergence terms, and water level gradient
  - Fully-Implicit: water level gradient term
- By treating the “fast” pressure gradient term implicitly, the time step limitation based on the wave celerity can be removed
- Both DWE and SWE use **Finite-Difference** and **Finite-Volume** Methods
- Time integration: **Finite-Difference**
- Continuity Equation: **Finite-Volume**
- Momentum Equation: **Finite-Difference** (no control volume)



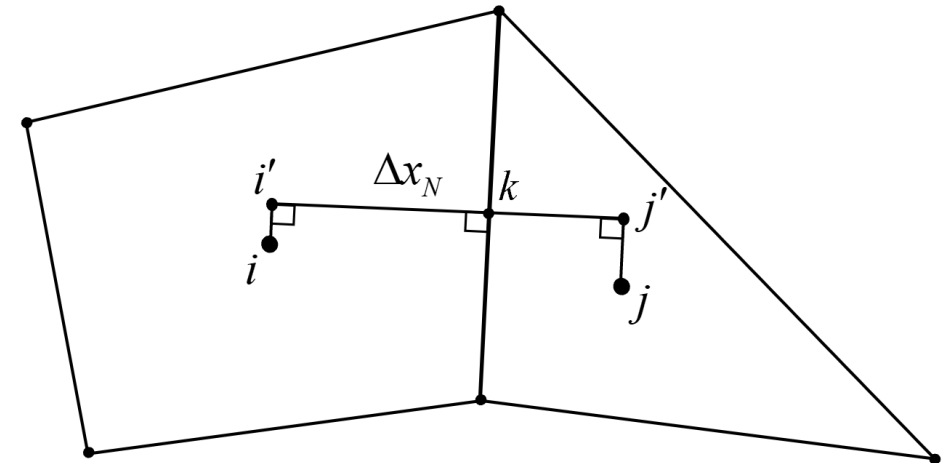
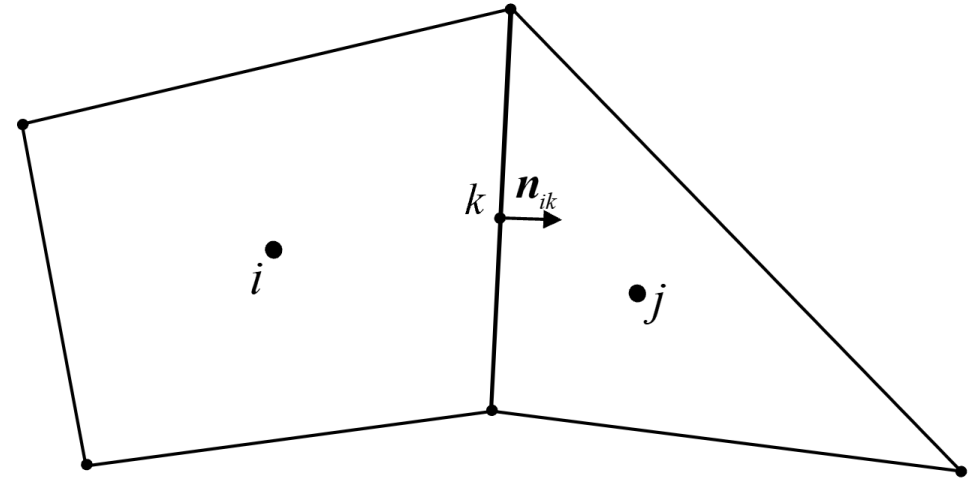
# Face Water Surface Gradient



- Face-Normal Gradient

$$\nabla H \cdot \mathbf{n}_{ik} = \frac{\partial H}{\partial N} \approx \frac{H_j - H_i}{\Delta x_N}$$

- Uses **Cell Centroids** and **NOT the Computation Points**
- Future versions may include non-orthogonal
- Compact two-point stencil is computationally efficient and robust





# Momentum Conservation

- Momentum conservation is directionally invariant
- Only “face-normal” component is needed at faces so

$$\frac{\partial u_N}{\partial t} + (\mathbf{V} \cdot \nabla) u_N - f u_T = -g \frac{\partial H}{\partial N} + \frac{1}{h} \nabla \cdot (\mathbf{v}_t h \nabla u_N) - c_f u_N + \frac{\tau_{s,N}}{\rho h}$$

where  $u_N$  is the velocity in the  $N$  direction

# Face-Tangential Velocity

- Tangential velocities are computed on left and right of face with a Least-squares Formulation

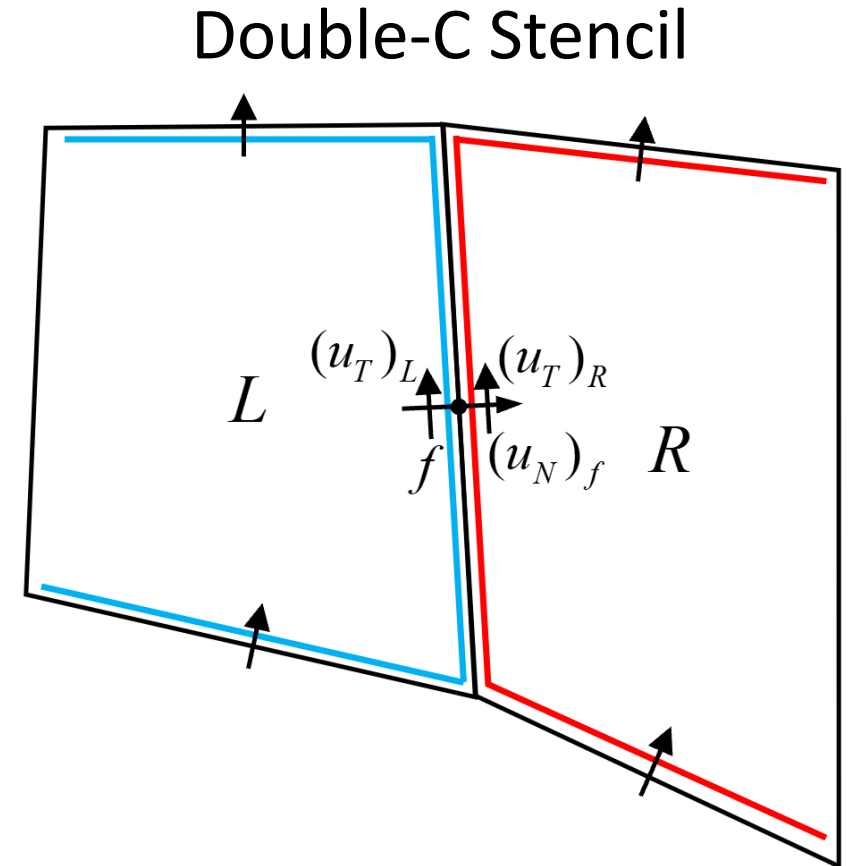
$$S_R = \sum_{k \in R} (\mathbf{V}_R \cdot \mathbf{n}_k - (u_N)_k)^2 \quad S_L = \sum_{k \in L} (\mathbf{V}_L \cdot \mathbf{n}_k - (u_N)_k)^2$$

- Of the left and right reconstructed velocities, only the tangential component is used, because the normal component is known

$$(u_T)_R = \mathbf{V}_R \cdot \mathbf{t}_f \quad (u_T)_L = \mathbf{V}_L \cdot \mathbf{t}_f$$

- Average face-tangential velocity computed as

$$(u_T)_f = \frac{(u_T)_R + (u_T)_L}{2}$$



# Discretization

- Cell Velocity Gradient (x-direction)

- Gauss' Divergence Theorem

$$\nabla u_i = \frac{1}{A_i} \int_A \nabla u dA = \frac{1}{A_i} \oint_L u \mathbf{n} dL = \frac{1}{A_i} \sum_{k \in i} u_k \mathbf{n}_{ik} L_k$$

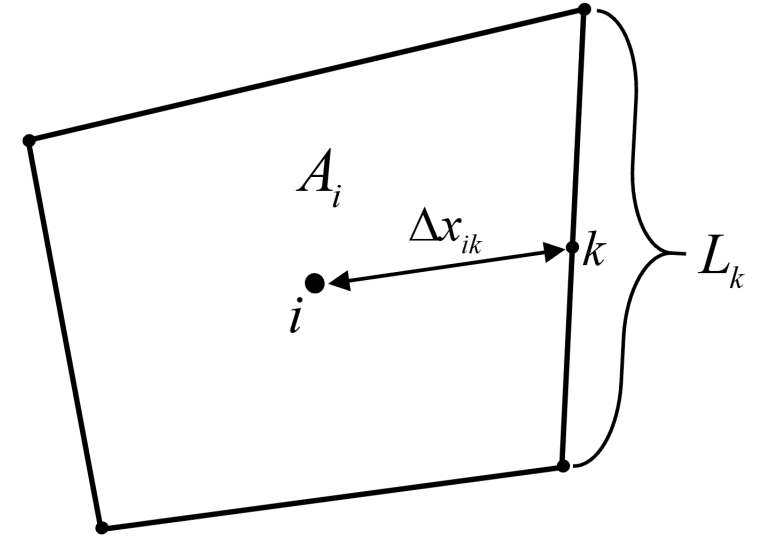
- Needed for turbulence modeling

- Cell Velocity

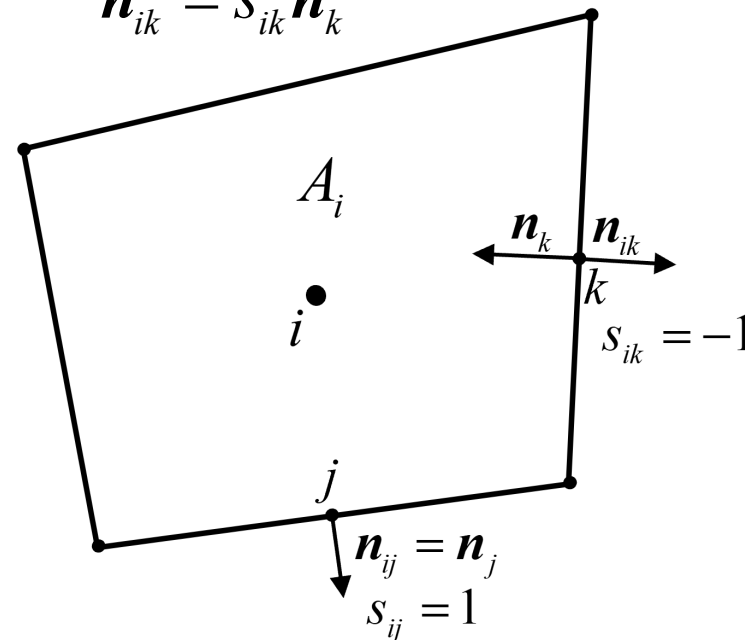
- Perot's Method

$$\mathbf{V}_i = \frac{1}{A_i} \sum_{k \in i} \Delta x_{ik} L_k \mathbf{n}_k (u_N)_k$$

- Needed for the conservative form of the mixing term and for Eulerian advection



$$\mathbf{n}_{ik} = S_{ik} \mathbf{n}_k$$





# Discretization: Laplacian

- Node Laplacian

$$(\nabla^2 V)_j = [\nabla \cdot (\nabla V)]_j \approx \sum_i d_i (\nabla V)_i$$

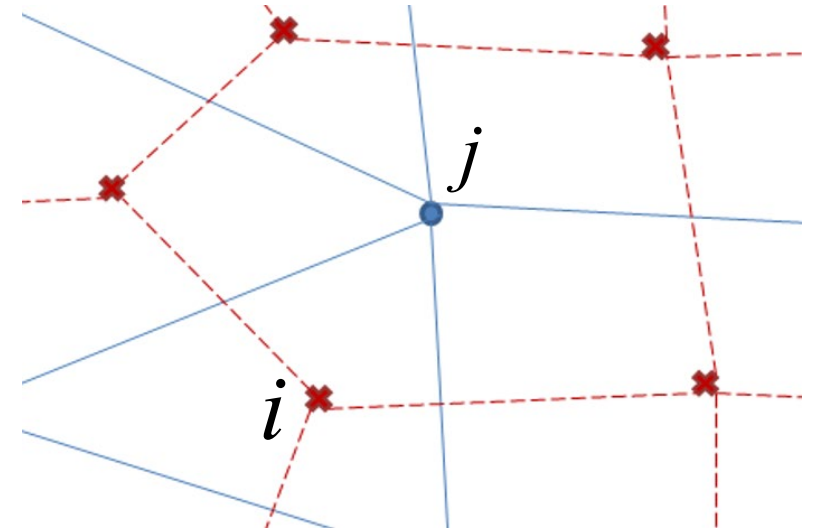
$i$  : Cells

$$(\nabla V)_i = \sum_k c_k V_k$$

$j$  : Nodes

$k$  : Faces

- Used only by non-conservative turbulence



# Backtracking

1. Interpolate node velocities from faces
2. Set starting location and time as  $A = f$ ,  $T_R = \Delta t$
3. From starting location and velocity, find location  $B$
4. Compute time to location  $B$ :  $T_B = (\mathbf{x}_A - \mathbf{x}_B) \mathbf{V}_A^{-1}$
5. Interpolate velocity at location  $B$ :  $\mathbf{V}_B = w_{n1} \mathbf{V}_{n1} + w_{n2} \mathbf{V}_{n2}$   
if  $T_B > T_R$

6. Set  $A = B$ ,  $T_R = T_R - T_B$ , and go to step 3  
else

7. Find location  $X$  as  $\mathbf{x}_X = \mathbf{x}_f - T_R \mathbf{V}_A$
8. Interpolate velocity vector at  $X$

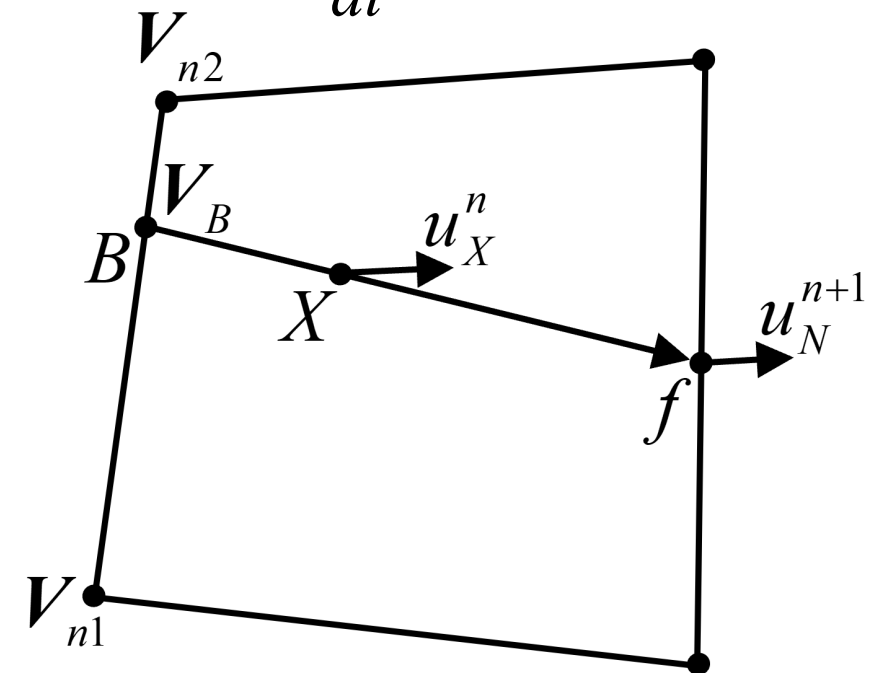
$$\mathbf{V}_X = T_B^{-1} \left[ T_R \mathbf{V}_B + (T_B - T_R) \mathbf{V}_A \right]$$

9. Compute advective velocity

$$u_X = \mathbf{n}_f \cdot \mathbf{V}_X$$

$$\frac{\partial u_N}{\partial t} + (\mathbf{V} \cdot \nabla) u_N = \frac{Du_N}{Dt} \approx \frac{u_N^{n+1} - u_X^n}{\Delta t}$$

$$\frac{d\mathbf{x}_P}{dt} = \mathbf{V}(\mathbf{x}, t)$$





# Fractional Step Method (ELM only)

- Coriolis Term approximated as

$$f\mathbf{k} \times \mathbf{V} \approx \begin{pmatrix} f \left[ (1-\theta)fv_X^n + \theta v^{n+1} \right] \\ -f \left[ (1-\theta)fu_X^n + \theta u^{n+1} \right] \end{pmatrix}$$

where

$f$  : Coriolis Parameter

$\theta$  : Implicit weighting factor

$\mathbf{k}$  : Unit vector in the vertical direction

$\mathbf{V} = (u, v)^T$  : Velocity at face

$\mathbf{V}_X = (u_X, v_X)^T$  : Velocity at face at location  $X$

- First (Coriolis) Step

$$\begin{pmatrix} 1 & \theta\Delta tf \\ \theta\Delta tf & 1 \end{pmatrix} \begin{pmatrix} u^* \\ v^* \end{pmatrix} = \begin{pmatrix} u_X^n + (1-\theta)\Delta tfv_X^n \\ v_X^n + (1-\theta)\Delta tfu_X^n \end{pmatrix} \quad \mathbf{V}^* = \begin{pmatrix} u^* \\ v^* \end{pmatrix}$$

- Second Step includes all other terms





# Eulerian-Lagrangian Momentum Equation

- Semi-discrete form (2<sup>nd</sup> Fractional Step)

$$\frac{u_N^{n+1} - u_N^*}{\Delta t} = -g \frac{\partial H^{n+\theta}}{\partial N} + \left[ \frac{1}{h} \nabla \cdot (\mathbf{v}_t h^n \nabla u_N) \right]_X^n - c_f u_N^{n+1} + \frac{\tau_{s,N}}{\rho h_f^n}$$

where

$$H^{n+\theta} = (1 - \theta)H^n + \theta H^{n+1}$$

$$u_N^* = \mathbf{V}^* \cdot \mathbf{n}_f$$

- Velocity  $\mathbf{V}^*$  includes Coriolis
- Mixing term is interpolated at backtracking location  $X$  and based on previous time step velocity field
- Friction term is semi-implicit



# Eulerian Momentum Equation

- Semi-discrete form

$$\frac{u_N^{n+1} - u_N^n}{\Delta t} + (\mathbf{V}^n \cdot \nabla) u_N^n - f u_T^n = -g \frac{\partial H^{n+\theta}}{\partial N} + \left[ \frac{1}{h} \nabla \cdot (\mathbf{v}_t h \nabla u_N) \right]_f^n - c_f u_N^{n+1} + \frac{\tau_{s,N}}{\rho h_f^n}$$

where

$$H^{n+\theta} = (1 - \theta) H^n + \theta H^{n+1} \quad \bar{h}_f = \alpha_f^L h_L + \alpha_f^R h_R$$

- Coriolis term computed at face  $f$  and is explicit
- No fractional step method like ELM solver
- Mixing term is computed at face  $f$  and is explicit
- Friction and pressure gradient terms are semi-implicit



# Discretization: Eulerian Advection

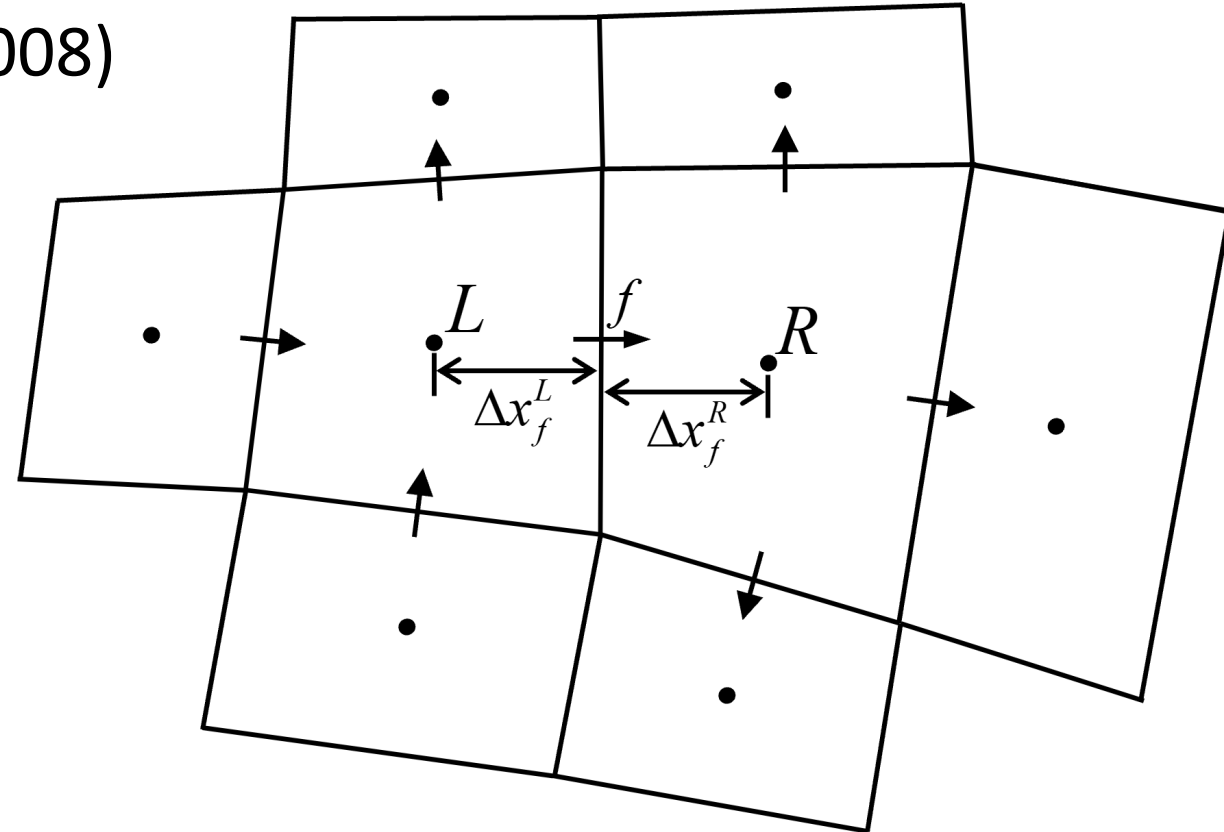
- Approach from Kramer and Stelling (2008)

$$(\mathbf{V} \cdot \nabla) u_N \approx \frac{\alpha_f^L}{\bar{h}_f A_L} \sum_{k \in L} s_{Lk} Q_k \left[ \mathbf{V}_k^u \cdot \mathbf{n}_f - (u_N)_f \right]$$

$$+ \frac{\alpha_f^R}{\bar{h}_f A_R} \sum_{k \in R} s_{Rk} Q_k \left[ \mathbf{V}_k^u \cdot \mathbf{n}_f - (u_N)_f \right]$$

$$\alpha_f^L = \frac{\Delta x_f^L}{\Delta x_f^L + \Delta x_f^R} \quad \bar{h}_f = \alpha_f^L h_L + \alpha_f^R h_R$$

$$\alpha_f^R = 1 - \alpha_f^L$$



- Courant-Freidrichs-Lewy (CFL) Condition

$$C = \frac{U \Delta t}{\Delta x} \leq 1$$



# Discretization: Mixing Term

- Non-Conservative Form

$$\mathbf{v}_t \nabla^2 u_N \approx \mathbf{v}_{t,f}^n \left( \nabla^2 V \right)_X^n \cdot \mathbf{n}_f$$

- Conservative Form

$$\frac{1}{h} \nabla \cdot (\mathbf{v}_t h \nabla u_N) \approx \frac{\alpha_f^L}{\bar{h}_f A_L} \sum_{k \in L} A_k \mathbf{v}_{t,k} \frac{\mathbf{n}_f \cdot (V_j - V_L)}{\Delta x_{L,j}} + \frac{\alpha_f^R}{\bar{h}_f A_R} \sum_{k \in R} A_k \mathbf{v}_{t,k} \frac{\mathbf{n}_f \cdot (V_j - V_R)}{\Delta x_{R,j}}$$

- Discretization same for both ELM and EM solvers
- Approximate Stability Criteria for EM solver

$$\frac{v_t \Delta t}{\Delta x^2} \leq \frac{1}{2}$$

- ELM interpolates term to location X

$$\left[ \frac{1}{h} \nabla \cdot (\mathbf{v}_t h \nabla u_N) \right]_X^n$$



# Eulerian-Lagrangian vs. Eulerian SWE Solvers

- **ELM-SWE**
  - Only solver available in V5.0.7 and earlier
  - Default in V6.0
  - Not limited by Courant condition
  - Excellent stability
  - Can have momentum conservation problems around shocks or where the flow changes rapidly
- **EM-SWE**
  - New to V6.0 as an option
  - Limited to Courant less than 1.0
  - Good Stability
  - Improved momentum conservation for all flow conditions

Strength/Feature/Capability	ELM-SWE	EM-SWE
Larger Time Step	X	
Best Stability	X	
Courant Stability Criteria		X
Diffusion Stability Criteria		X
Computational Speed	X	
Wet/dry > 1 cell per time step	X	
Best Momentum Conservation		X
Non-Conservative Mixing		X
Conservative Mixing	X	X
Wind	X	X



# Subgrid Modeling

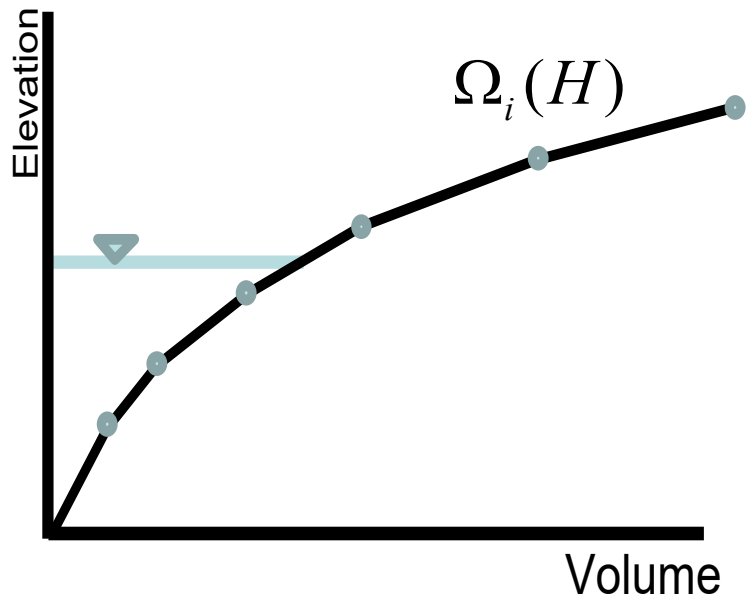
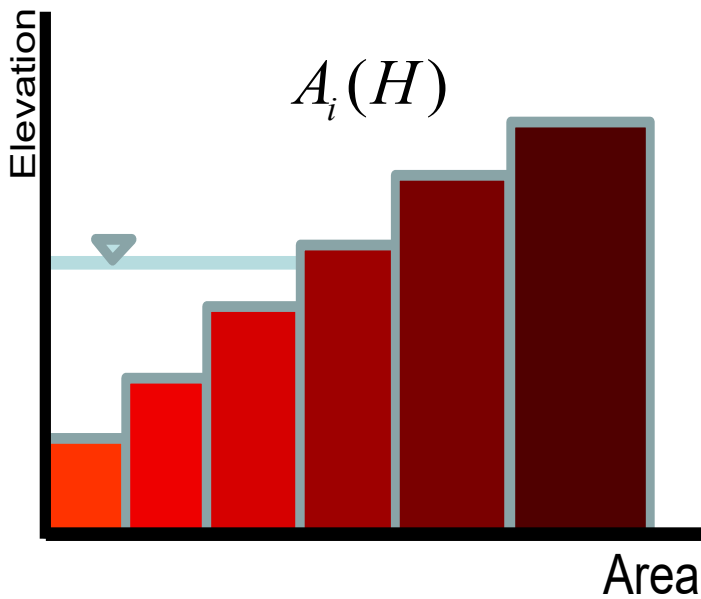
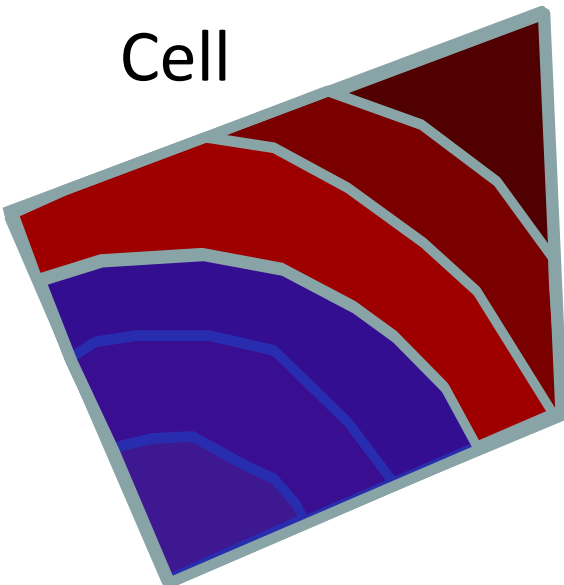


- Problem
  - Water levels usually vary much more smoothly than the terrain
  - Unfeasible to resolve every detail of the terrain with the computational mesh
- Approach
  - Utilize a grid resolution sufficient to resolve the hydraulics
  - Capture the details of the subgrid terrain through hydraulic properties tables



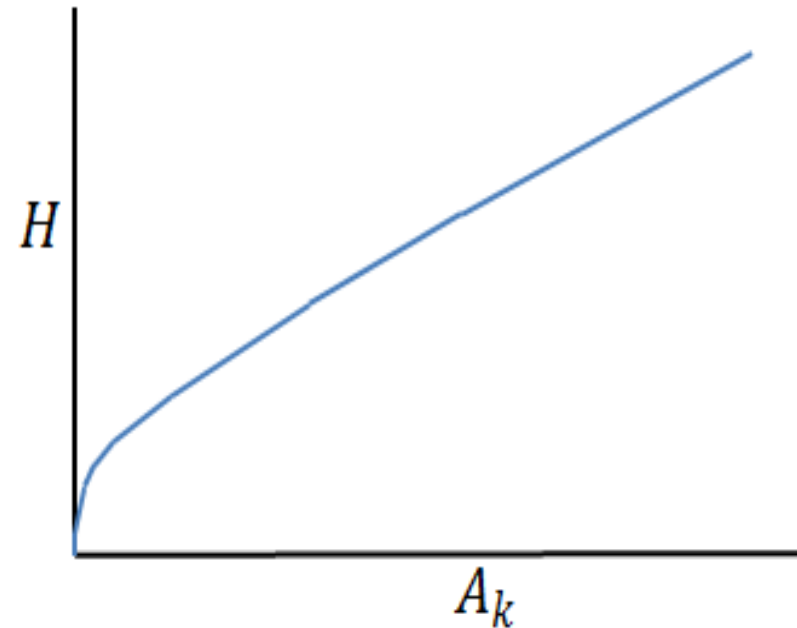
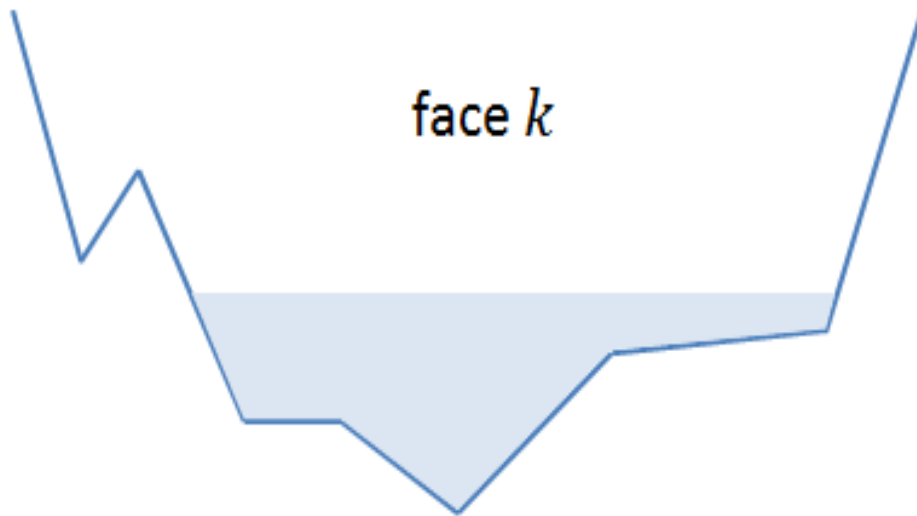


# Subgrid Bathymetry: Cells



# Subgrid Bathymetry: Faces

- Faces treated similar to cells
- Hydraulic property tables computed
  - Wetted length
  - Wetted Perimeter
  - Area







# Benefits of Subgrid Bathymetry





# Solution Procedure

For both DWE and SWE:

- The face-normal momentum equation is inserted into the continuity equation to obtain a non-linear and implicit equation for water levels
- For example, in the case of DWE

$$\frac{\Omega_i^{n+1} - \Omega_i^n}{\Delta t} - \sum_{k \in i} \alpha_k \left( \frac{H_j^{n+\theta} - H_i^{n+\theta}}{\Delta x_N} \right) = Q_i \quad \Omega_i = \Omega_i(H_i)$$

- Implicit weighting

$$H^{n+\theta} = (1 - \theta)H^n + \theta H^{n+1}$$

- Once the water levels are computed, the new time step face-normal velocities are computed



# Solution Procedure



- System of equations

$$\mathbf{\Omega} + \mathbf{\Psi}H = \mathbf{b}$$

- Algorithm

1. Compute Right-Hand-Side  $\mathbf{b}$ 
  - Contains explicit terms:  
advection, diffusion, wind, etc.
2. Outer Loop (Assembly and Updates)
  - Update linearized terms and variables  
including coefficient matrix  $\mathbf{\Psi}$
3. Inner Loop (Newton Iterations)

$H$  : Water level

$\mathbf{\Omega}$  : Water volume

$\mathbf{\Psi}$  : Coefficient matrix

$\mathbf{b}$  : Right-hand-side

$m$  : Iteration index

$\mathbf{A}$  : Diagonal matrix of  
cell wet surface areas

$$H^{m+1} = H^m - [\mathbf{\Psi} + \mathbf{A}^m]^{-1} (\mathbf{\Omega}^m + \mathbf{\Psi}H^m - \mathbf{b})$$



# Boundary Conditions

- **Stage Hydrograph.** Upstream or downstream
- **Flow Hydrograph.** Upstream or downstream. Local conveyance and velocities computed automatically.
- **Normal Depth BC.** At downstream boundaries.
- **Rating Curve BC.**
- **Wind.** Only for shallow-water equations.
- **Precipitation, evapotranspiration, and infiltration.** Included as sources and sinks in the continuity equation.
- 1D reaches and 2D areas can be connected
- Multiple 2D areas can be connected to each other
- 2D areas can be connected to 1D lateral structures such as levees to simulate levee breaches



# Computational Implementation



- Multiple 2D areas can be computed independently and simultaneously
- All solvers are can be run on multiple cores
- 2D solvers and parameters can be selected independently for each 2D area
- A partial grid solution keeps track of active portion of mesh and only computes the solution for active portion significantly reducing computational times.

# Thank You!

HEC-RAS Website:

<https://www.hec.usace.army.mil/software/hec-ras/>

Online Documentation:

<https://www.hec.usace.army.mil/confluence/rasdocs>